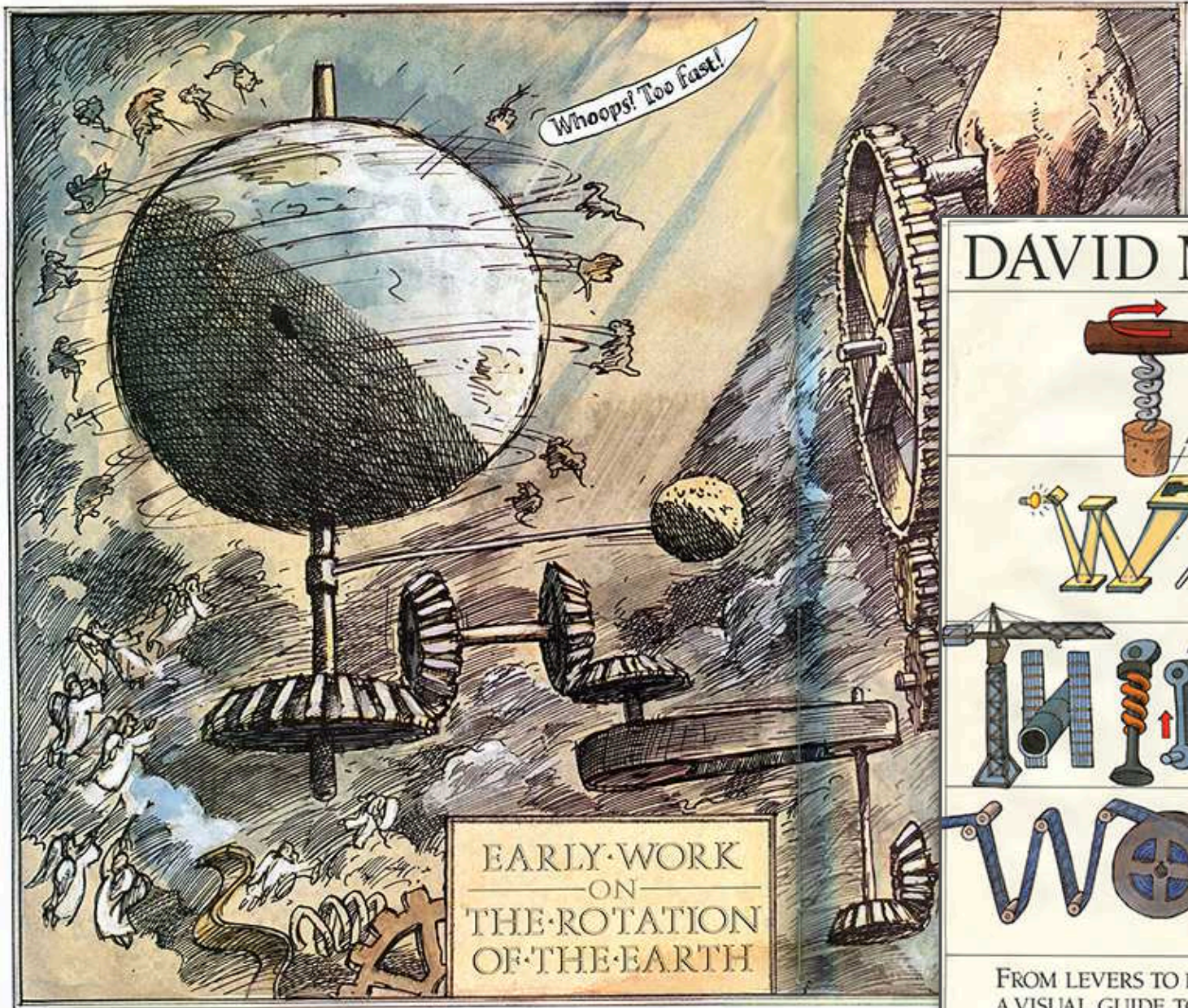
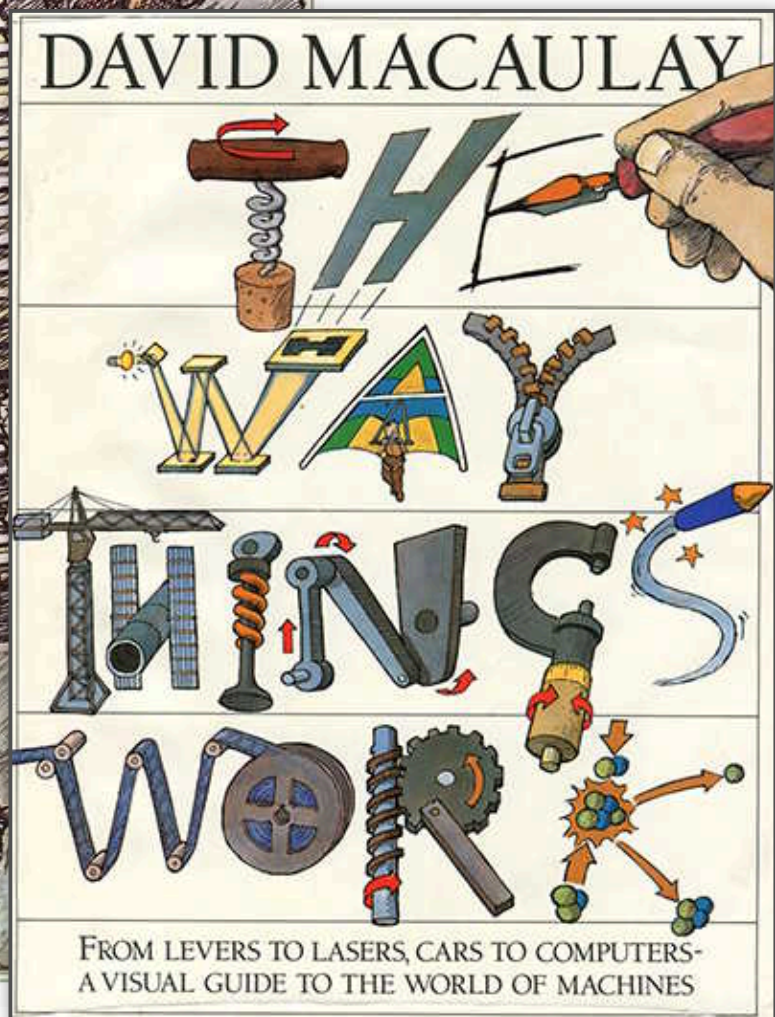


deep
earth

... oder wie Langsamkeit funktioniert



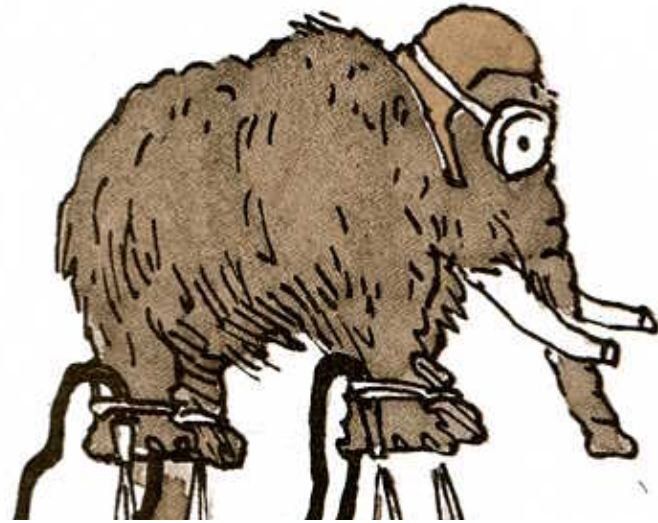
... and ...



... down there

brought to you

by:



Renée Heilbronner

Rüdiger Kilian

Anja Thust

Matěj Peč

Holger Stünitz, Tromsø

Jan Tullis, Brown Univ.

Brian Evans, MIT

Andreas Kronenberg, TAMU

Barbara John, Univ. Wyoming

SNF

Einleitung

Hauptteil

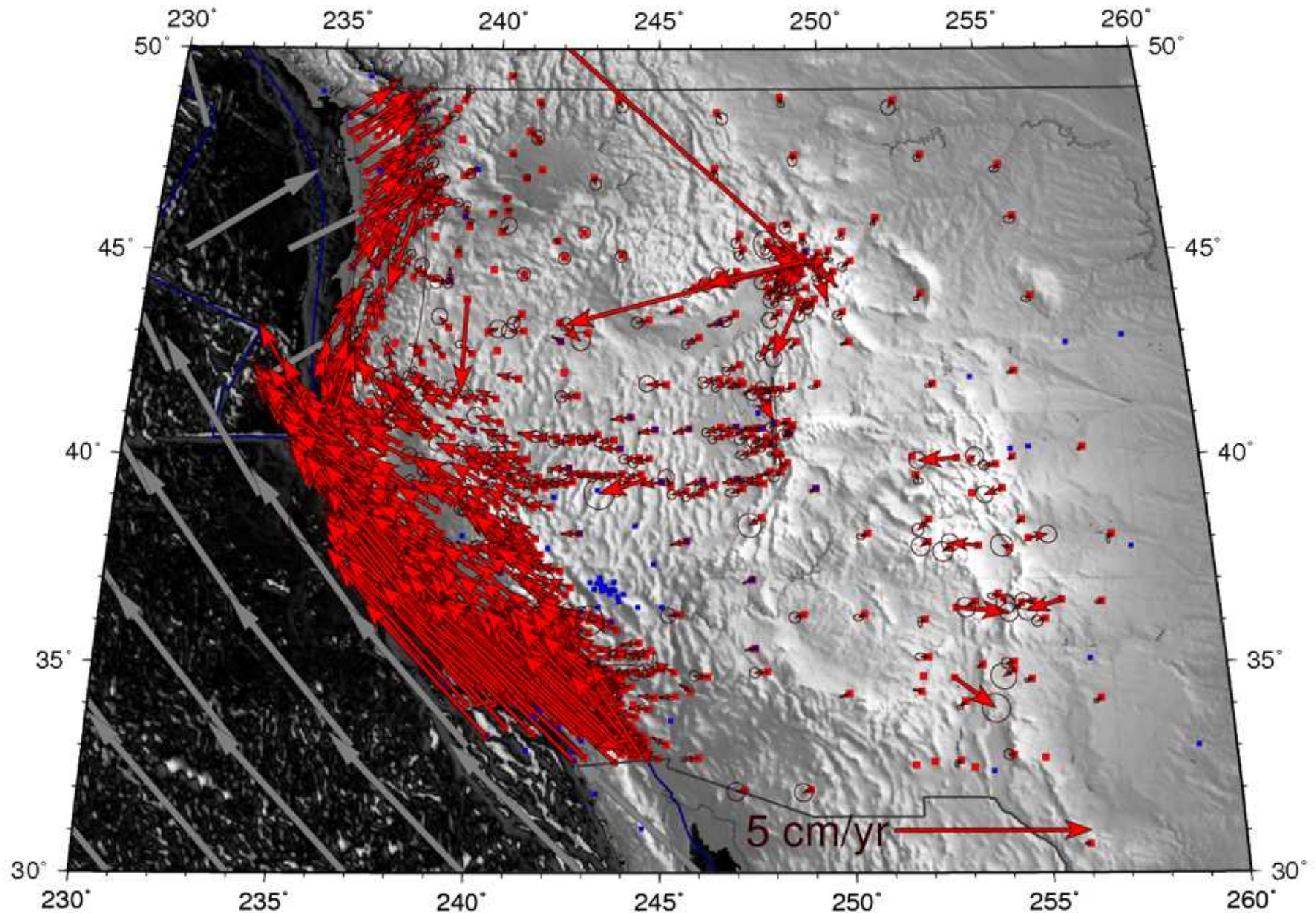
Schluss

PBO MIT GPS velocities

GMT 2008 Oct 4 09:03:49

Red: PBO (90% uncertainty, only data with $\sigma_v < 0.5$ cm/yr errors). Blue: UNAVCO GPSVEL1.0 sites. Grey: NUVEL-1A NAM fixed. Contact: twb@usc.edu

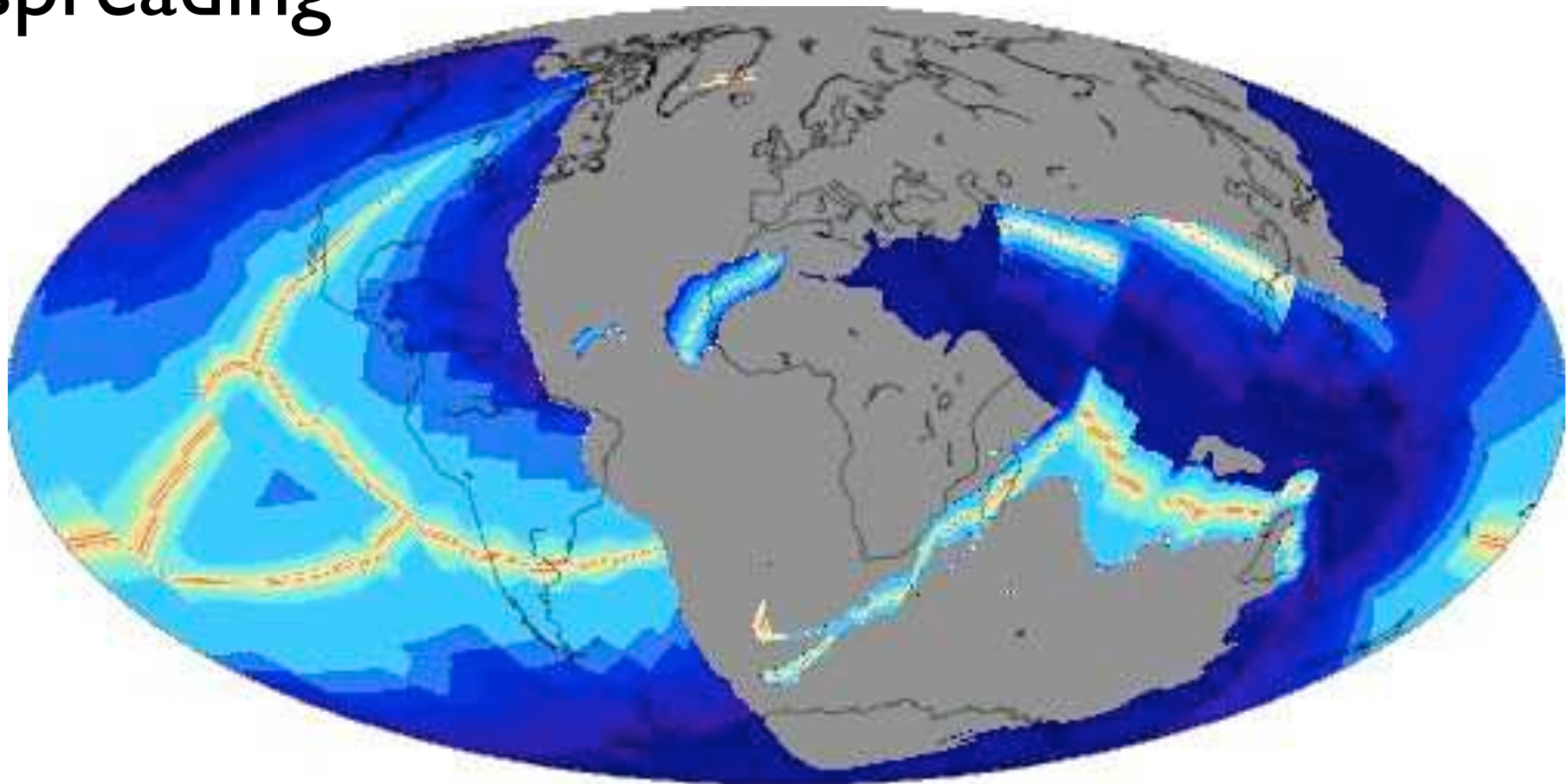
Release date: 2008-09-30 (20080930124552)



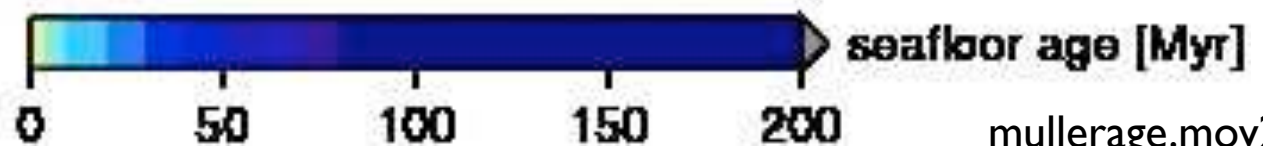
gemessene Langsamkeit plate motion

velocity = displacement / time
(in ms^{-1})

geological reconstruction from sea floor spreading



140 Ma



mullerage.mov2

berechnete Langsamkeit geological deformation

strain = displacement / length

strain rate = strain / time

(in s^{-1})

Beispiel: San Andreas Fault

Plattenbewegung pro Jahr

$$d = 10 \text{ cm}$$

Verformter Bereich

$$B = 100 \text{ km}$$

shear strain:

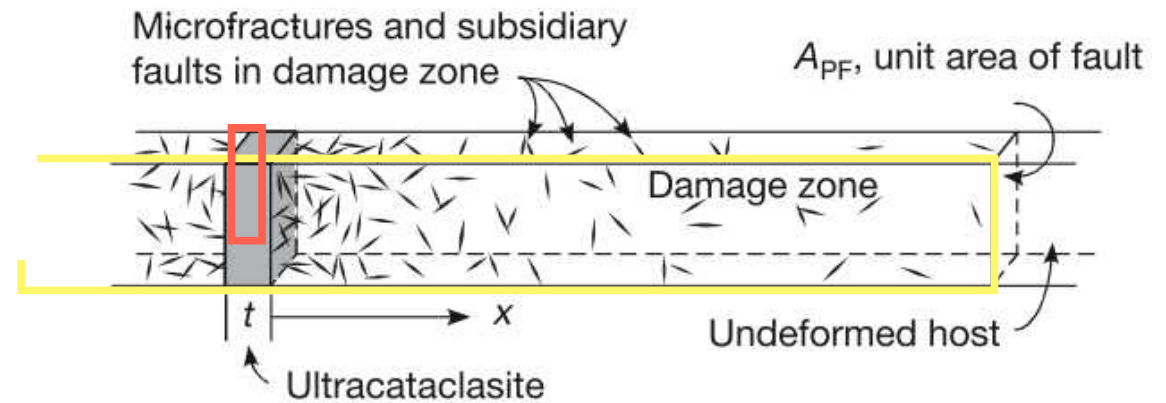
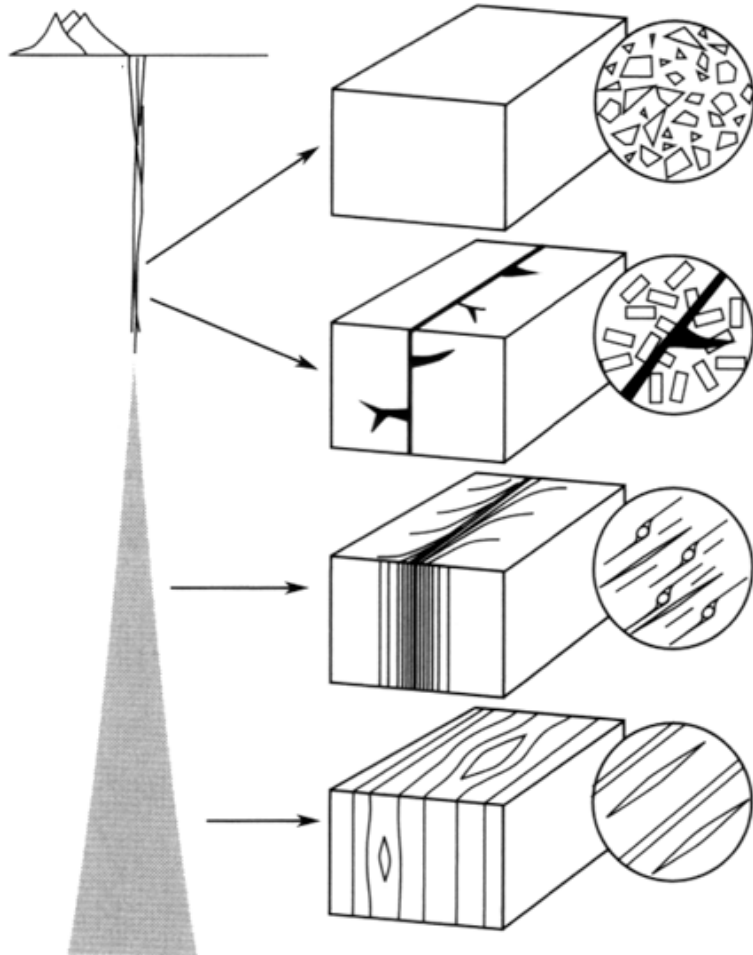
$$\gamma = d / B = 10^{-1} \text{ m} / 10^5 \text{ m} = 10^{-6}$$

strain rate:

$$\dot{\gamma} = \gamma / t = 10^{-6} / a = 10^{-6} / (\sim 3 \cdot 10^7 \text{ s}^1)$$

$$\dot{\gamma} \approx 10^{-14} \text{ s}^{-1}$$

Lokalisierung



$$100\text{km} \Rightarrow 100\text{m} = 10^3$$

$$\dot{\gamma} = 10^{-11} \text{ s}^{-1}$$

$$100\text{km} \Rightarrow 1\text{mm} = 10^8$$

$$\dot{\gamma} = 10^{-6} \text{ s}^{-1}$$

Lokalisierung

100km \Rightarrow 100m = 10^3

$$\dot{\gamma} = 10^{-11} \text{ s}^{-1}$$

1% der Zeit :

$$\dot{\gamma} = 10^{-9} \text{ s}^{-1}$$

fast !

100km \Rightarrow 1mm = 10^8

$$\dot{\gamma} = 10^{-6} \text{ s}^{-1}$$

1 Sekunde / Jahr:

$$\dot{\gamma} \approx 30 \text{ s}^{-1}$$

seismic !

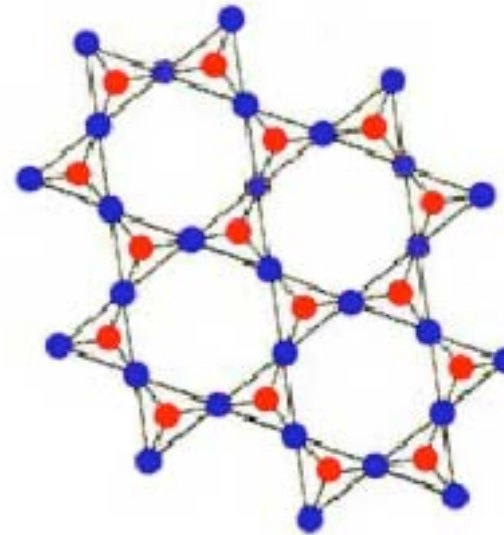
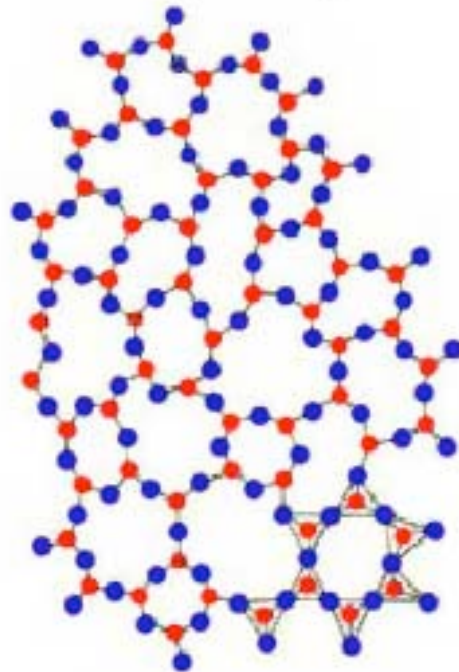
$$1 \text{ a} \approx 3 \cdot 10^7 \text{ s}$$

kontinuum-diskontinuum kristalline Festkörper

amorphes SiO_2
(Glas)

kristallines SiO_2
(Quarz)

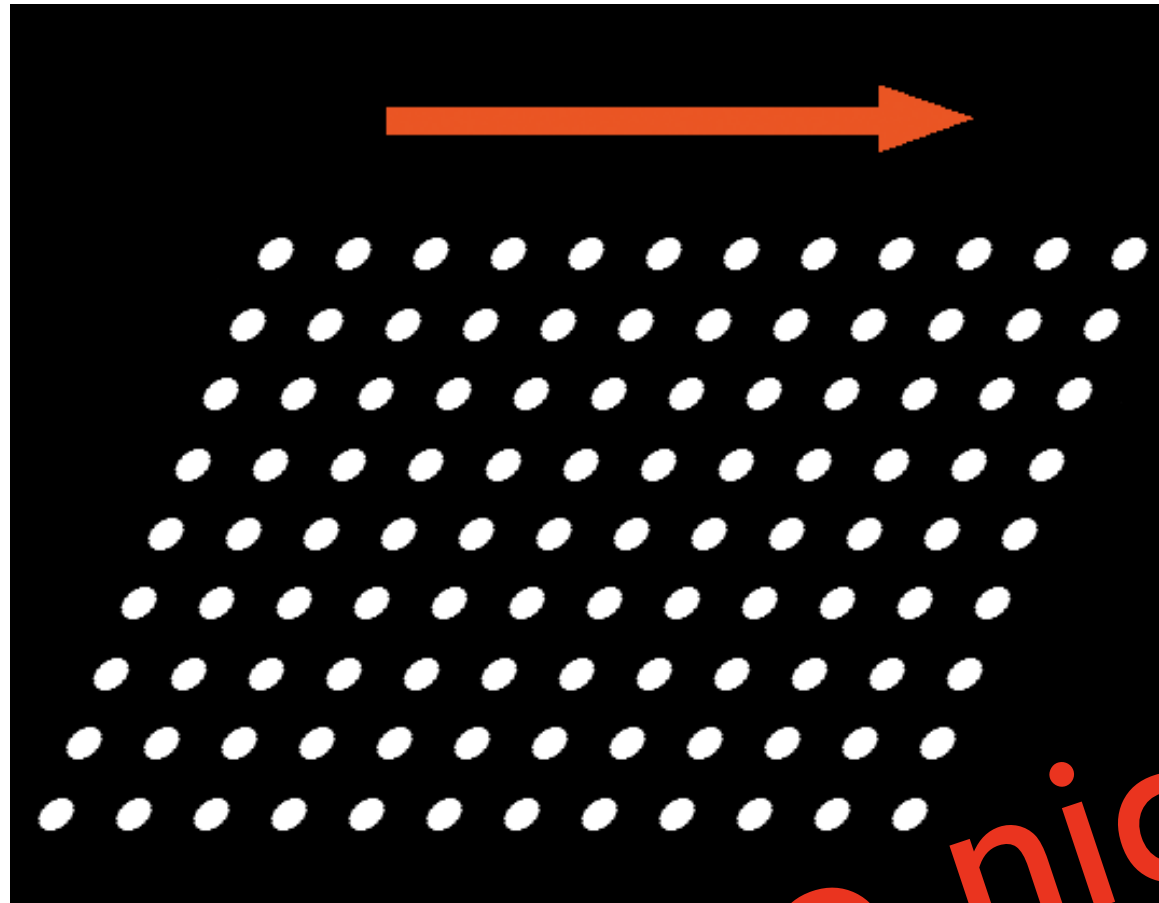
isotrop



O ● Si ●

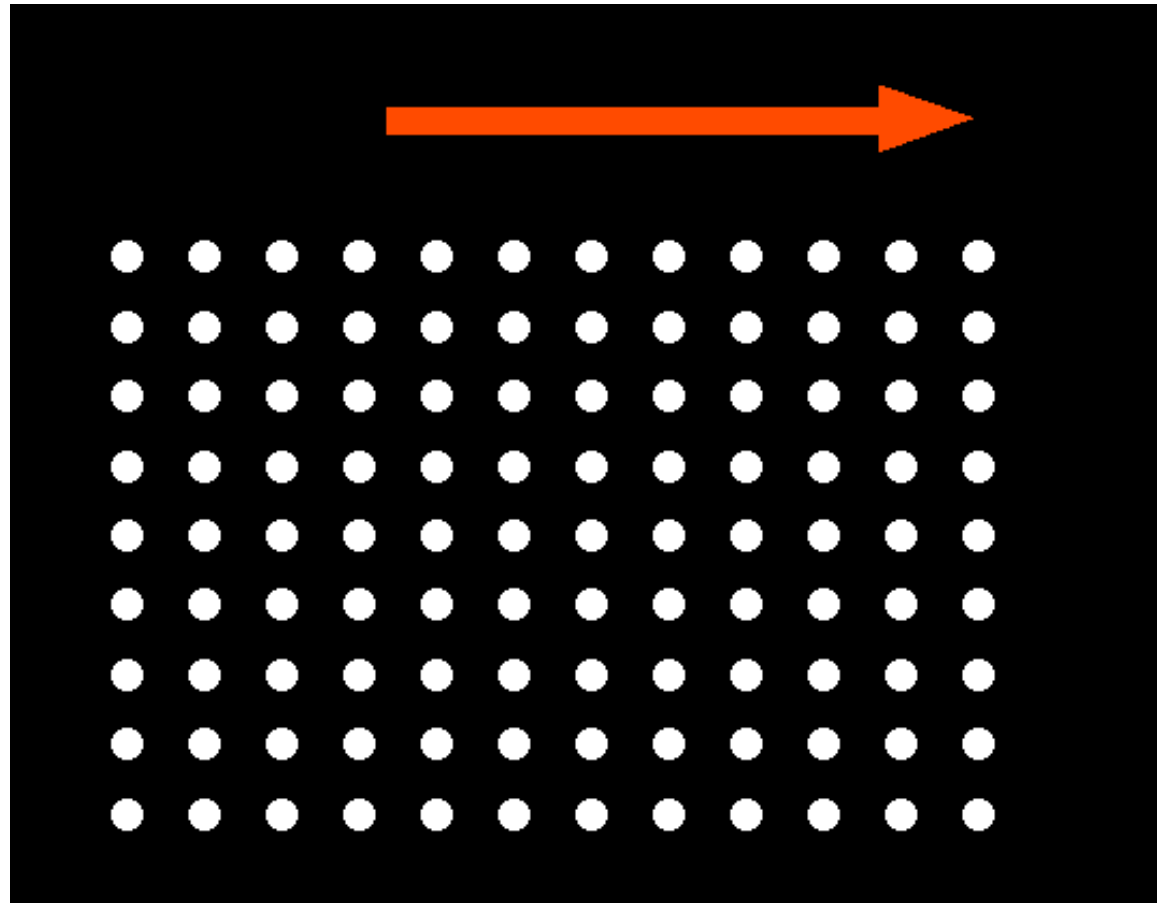
anisotrop

Deformation von Kristallen



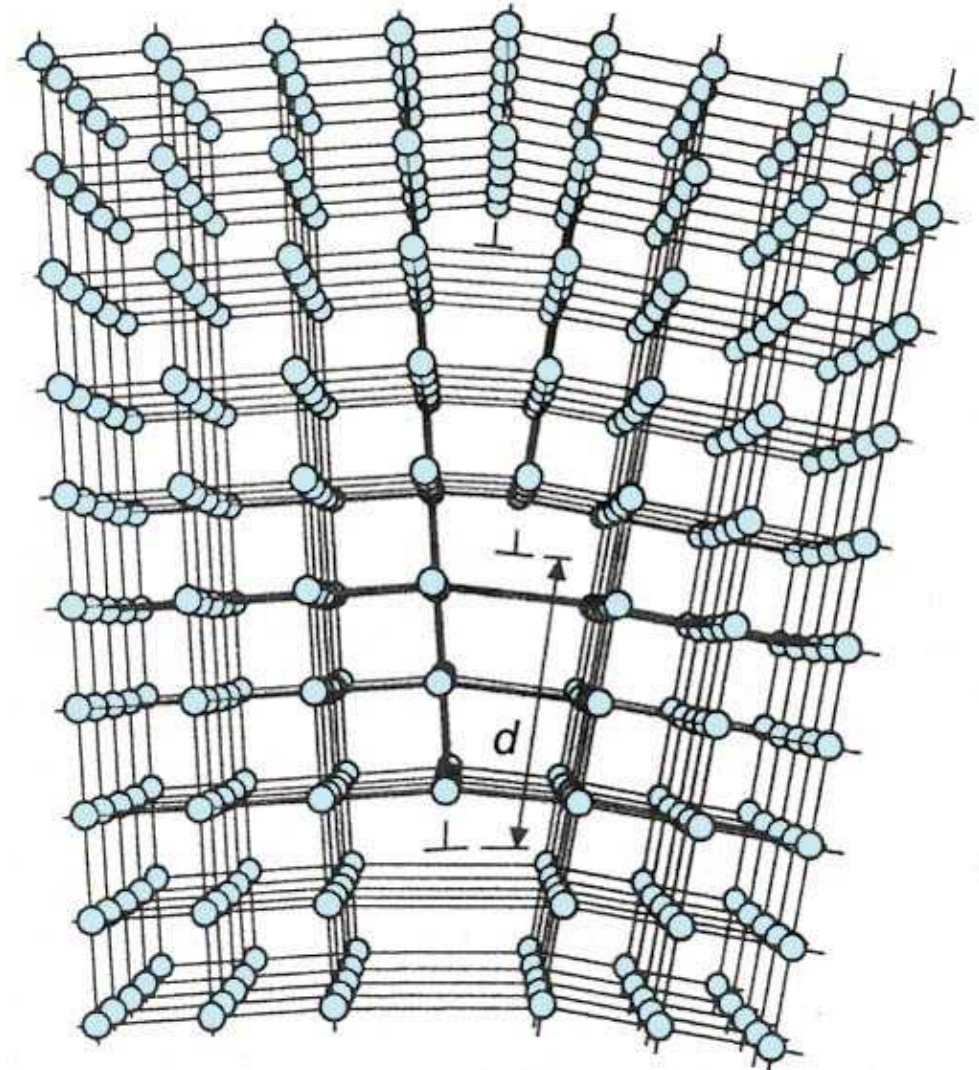
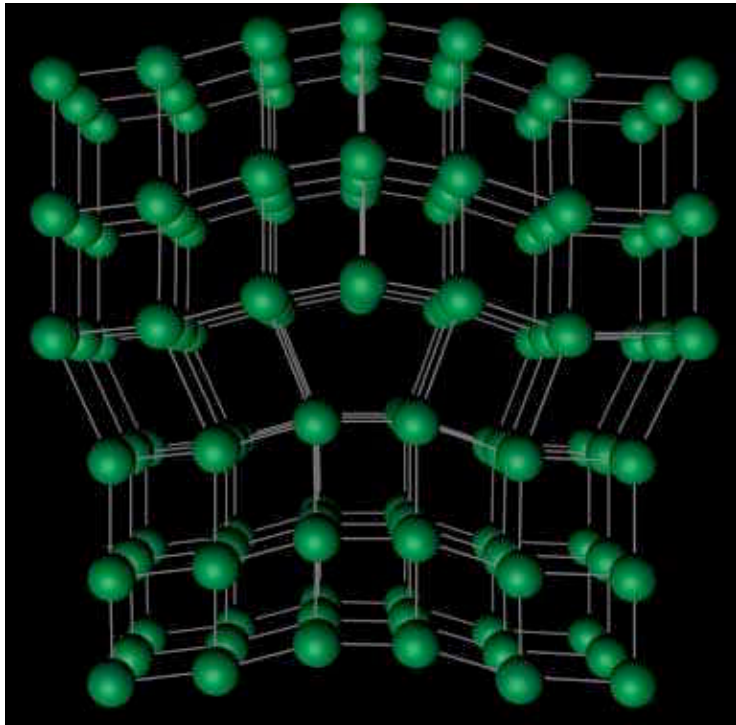
so nicht!

Versetzungen

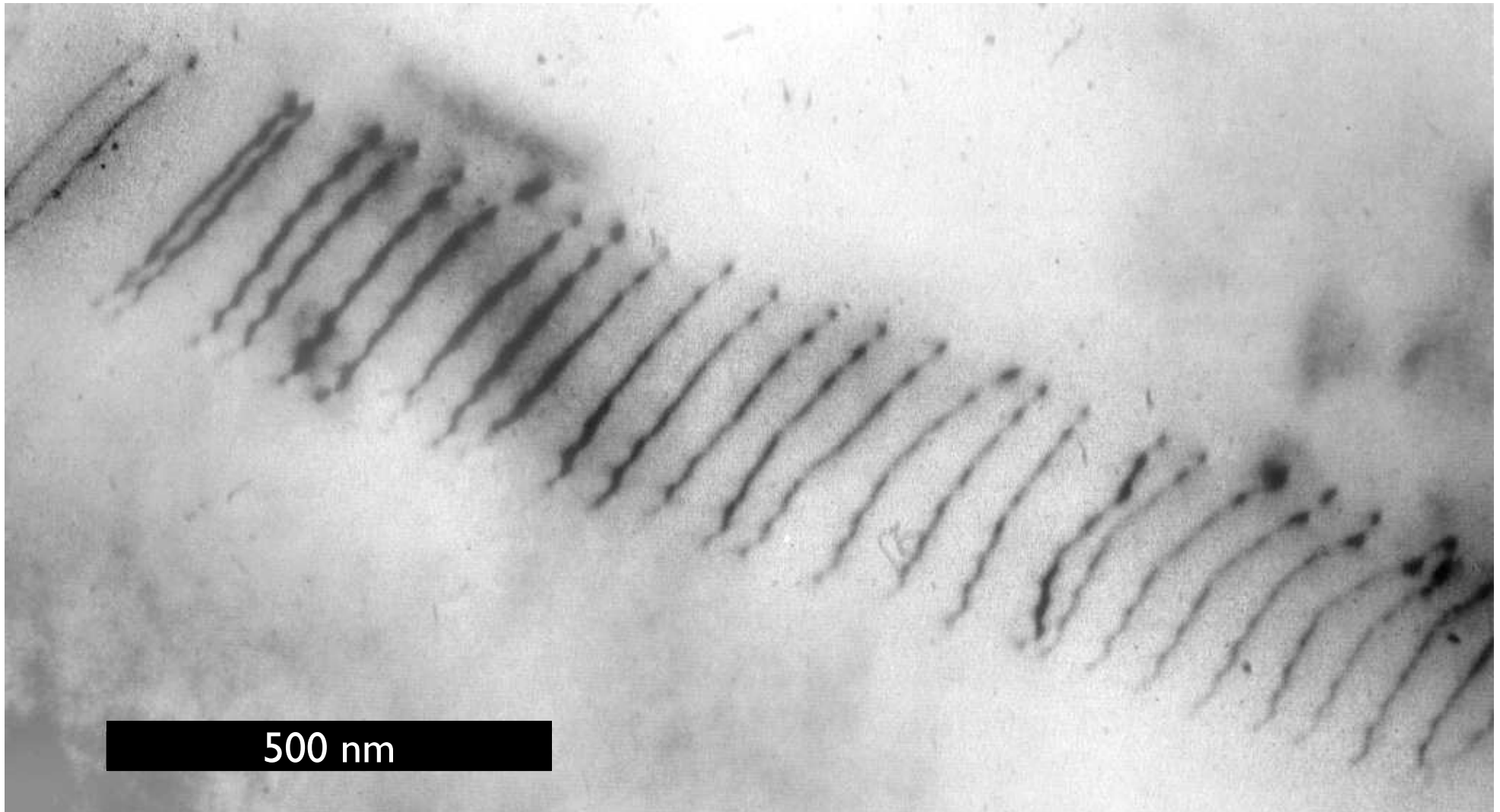


sondern so !

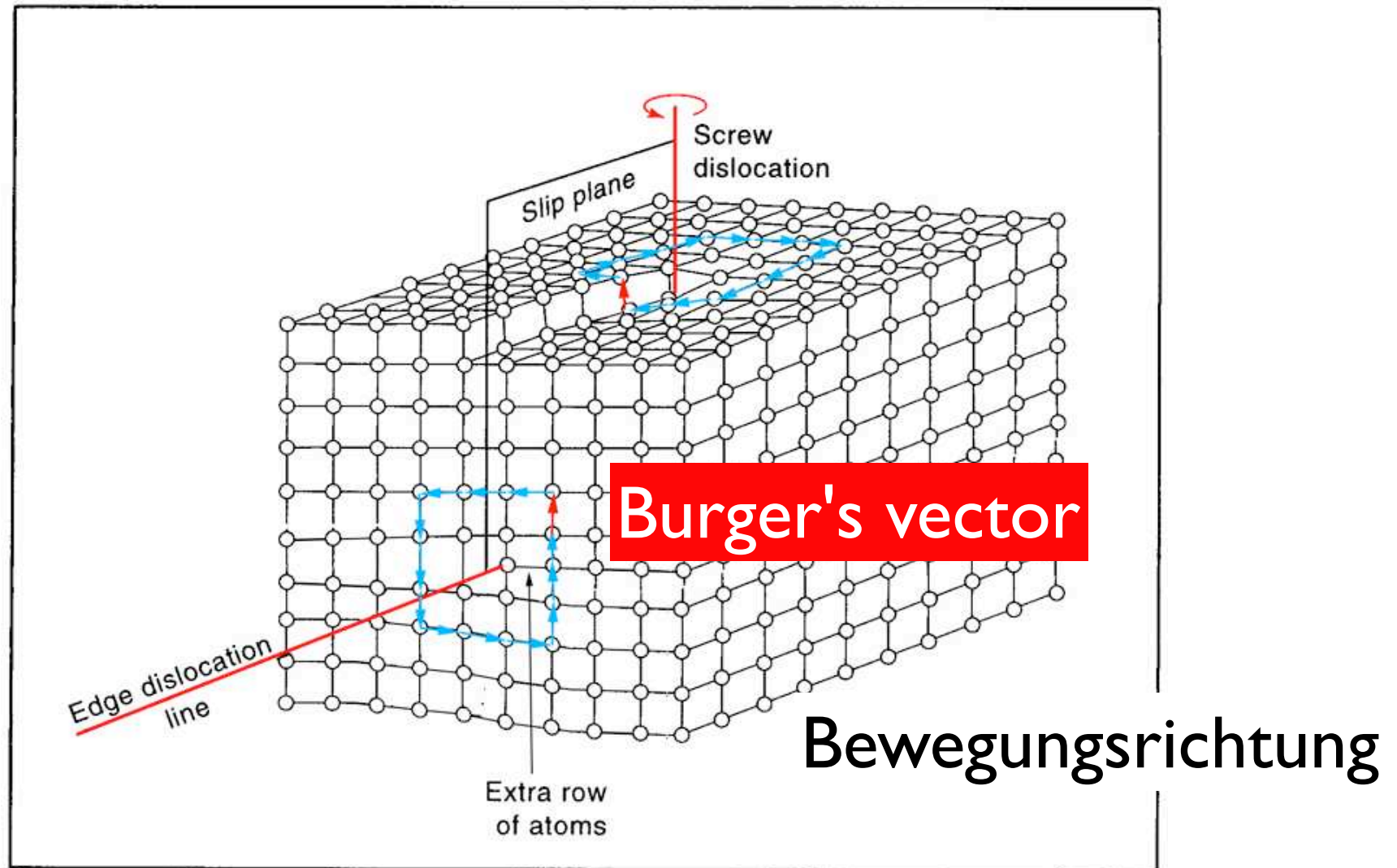
edge dislocation



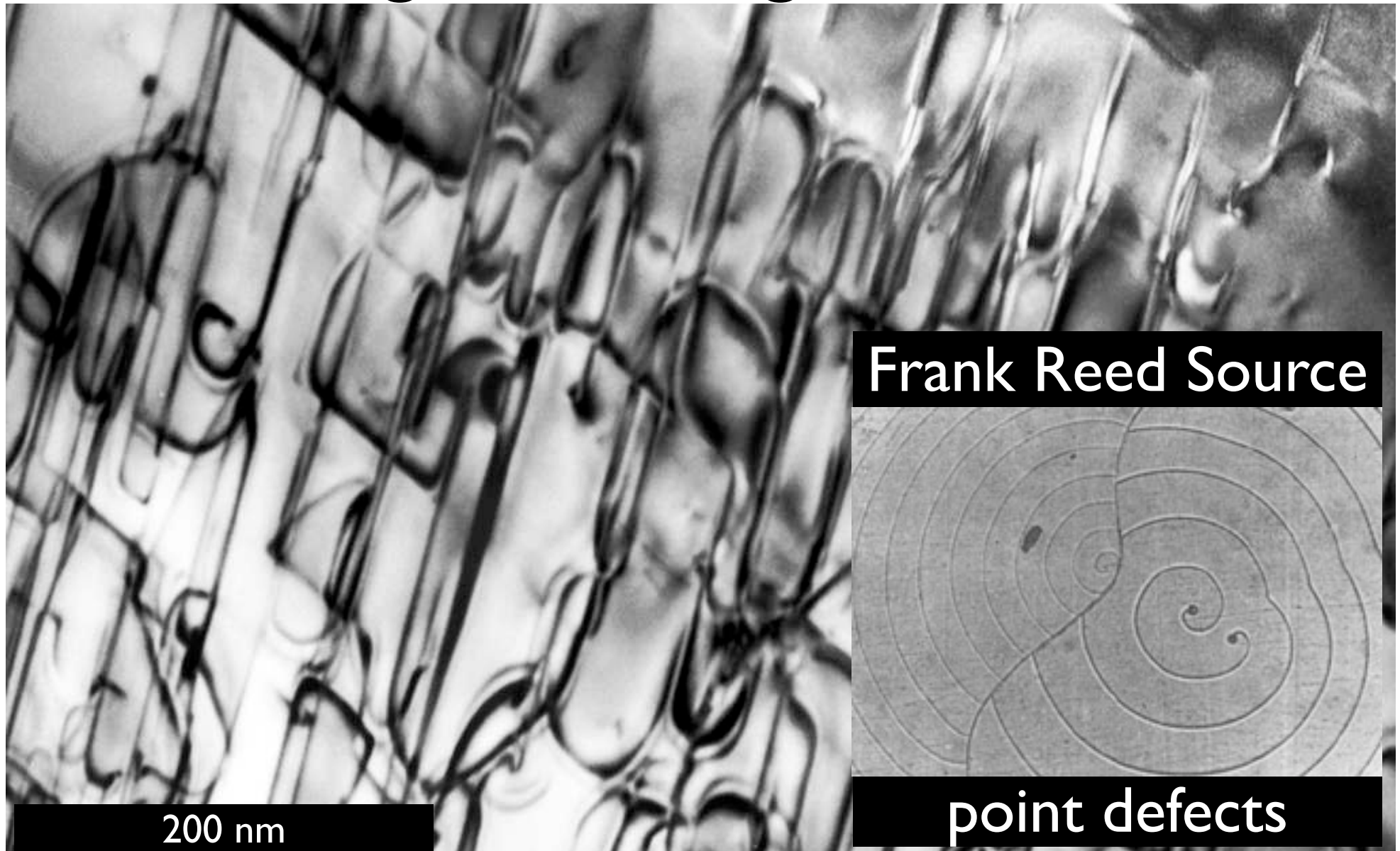
edge dislocation



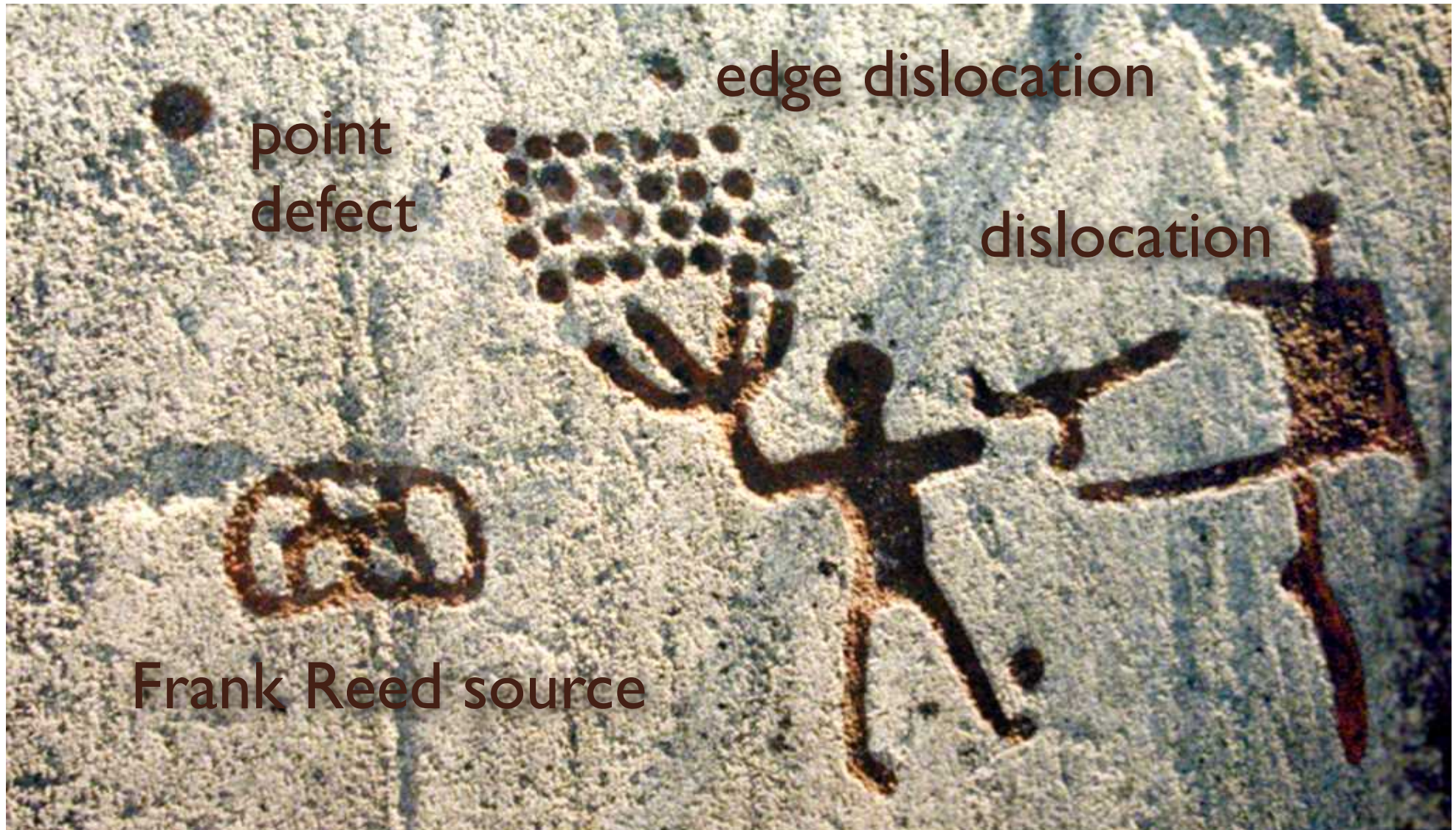
screw & edge dislocation



moving & creating dislocations



oldest dislocations



dislocation speed: supersonic

$$v = \sqrt{C_{ij} / \rho}$$

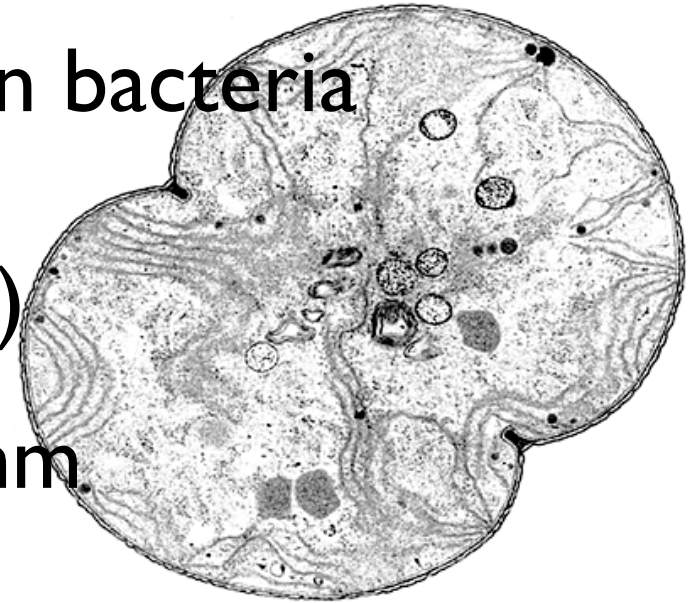
speed of sound in air = 355 ms⁻¹

in minerals ~ 6000 ms⁻¹

(in rubber ~ 60 ms⁻¹)

Vergleich:

The speed of DNA replication in bacteria
 ≈ 1000 nucleotides / s
(in humans ≈ 50 nucleotides / s)



Length of nucleotide unit: 0.33 nm

\Rightarrow replication speed (in ms^{-1}):

$$1000 \cdot 0.33 \text{ nm s}^{-1} = 330 \text{ pm s}^{-1} = 3.3 \cdot 10^{-10} \text{ ms}^{-1}$$

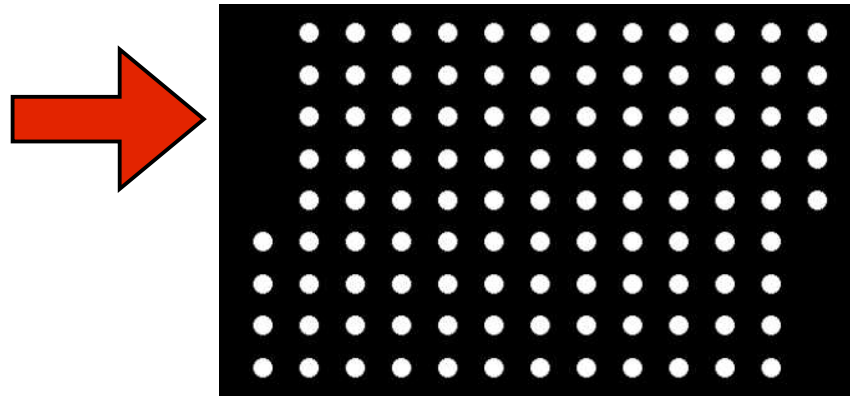
Assume dislocation speed (in rocks): $3.3 \cdot 10^3 \text{ ms}^{-1}$

$$\Rightarrow \text{GEO speed} \approx 10^{13} \cdot \text{BIO speed}$$

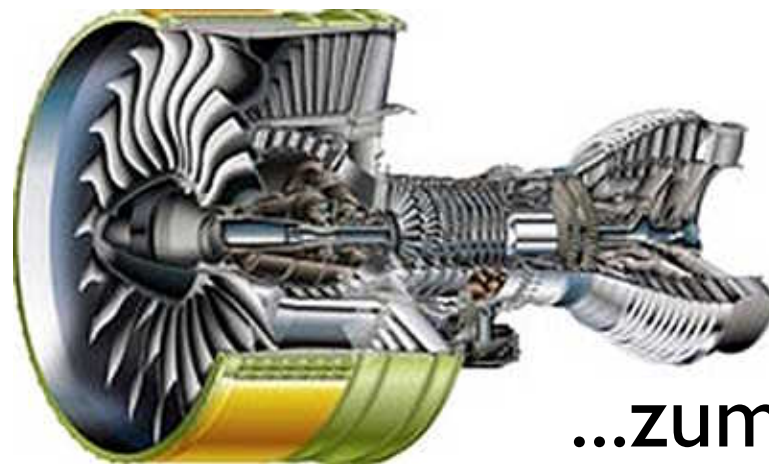
plastische Verformung von kristallinen Festkörpern (Aggregaten)

warum ergeben viele
schnelle Versetzungen
eine grosse Langsamkeit ?!

0. Es ist gar nicht so einfach Dislokationen zu erzeugen !



so einfach
geht's nicht

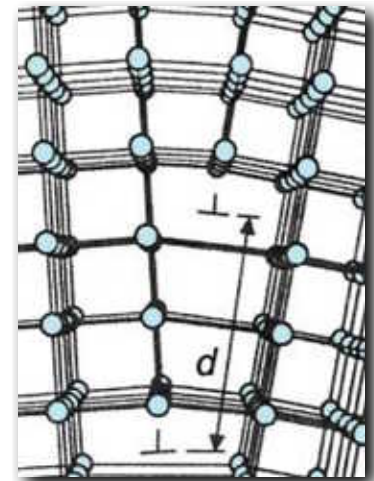


...zum Glück !

I. Nicht jede Richtung im Kristallgitter eignet sich zum Gleiten (**dislocation glide**), es gibt "harte" und "weiche" Gleitsysteme.

$$v = \sqrt{C_{ij} / \rho}$$

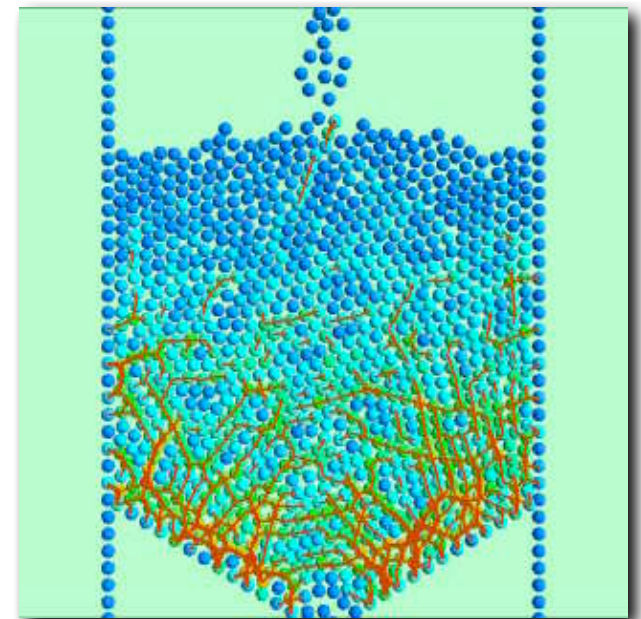
2. Dislokationen werden behindert (dislocation pile-up).
Geschwindigkeitsbestimmender Schritt \neq Dislokationsgleiten, sondern die Diffusion (Beseitigung) der Dislokationen.



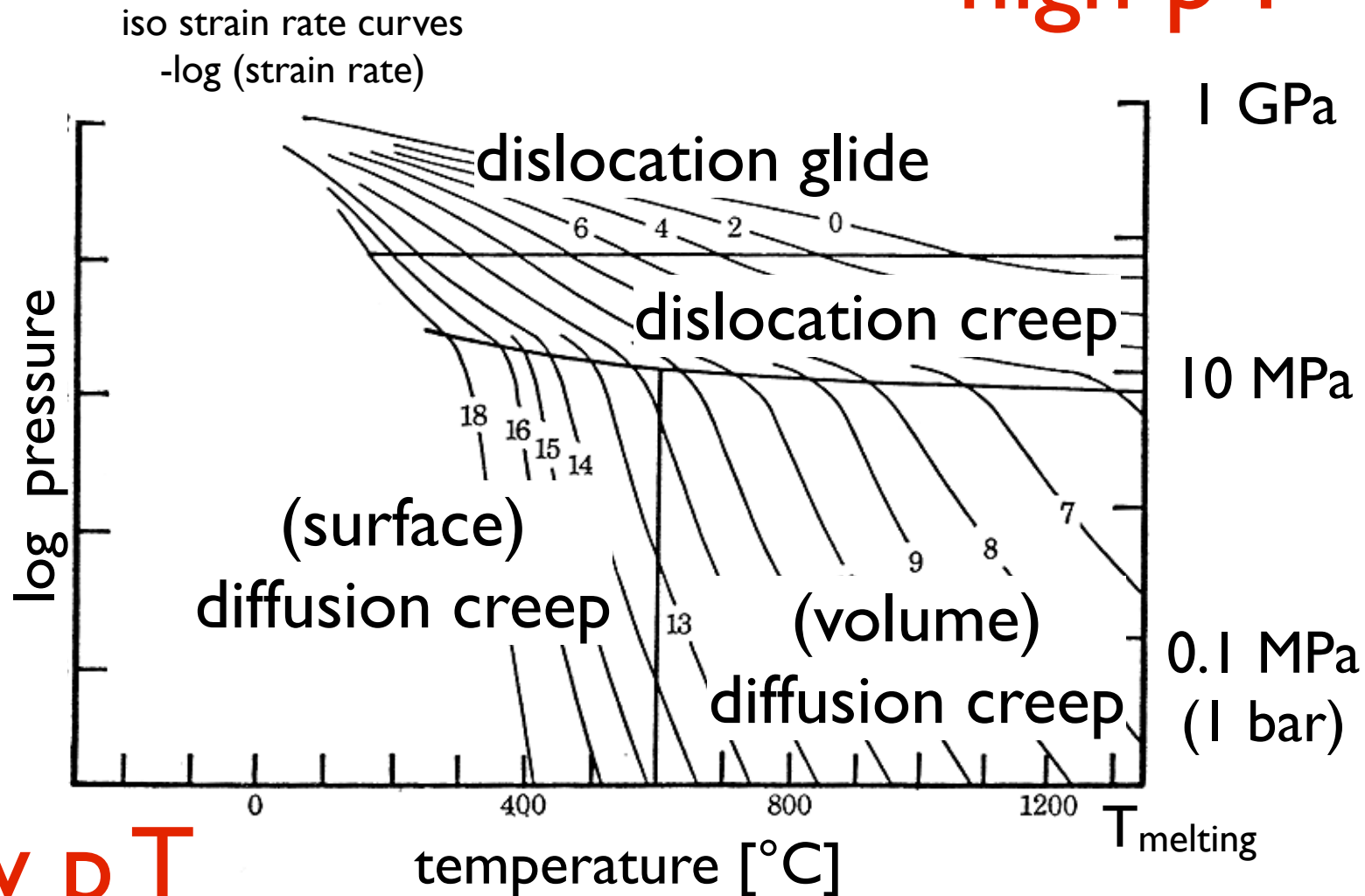
\Rightarrow **dislocation creep**

3. Verformung weicht aus auf schwächere Phasen, \pm kein Dislokationsgleiten, Korn gleitet oder diffundiert als Ganzes.

=> g boundary sliding,
granular flow,
diffusion creep



high p T



low p T

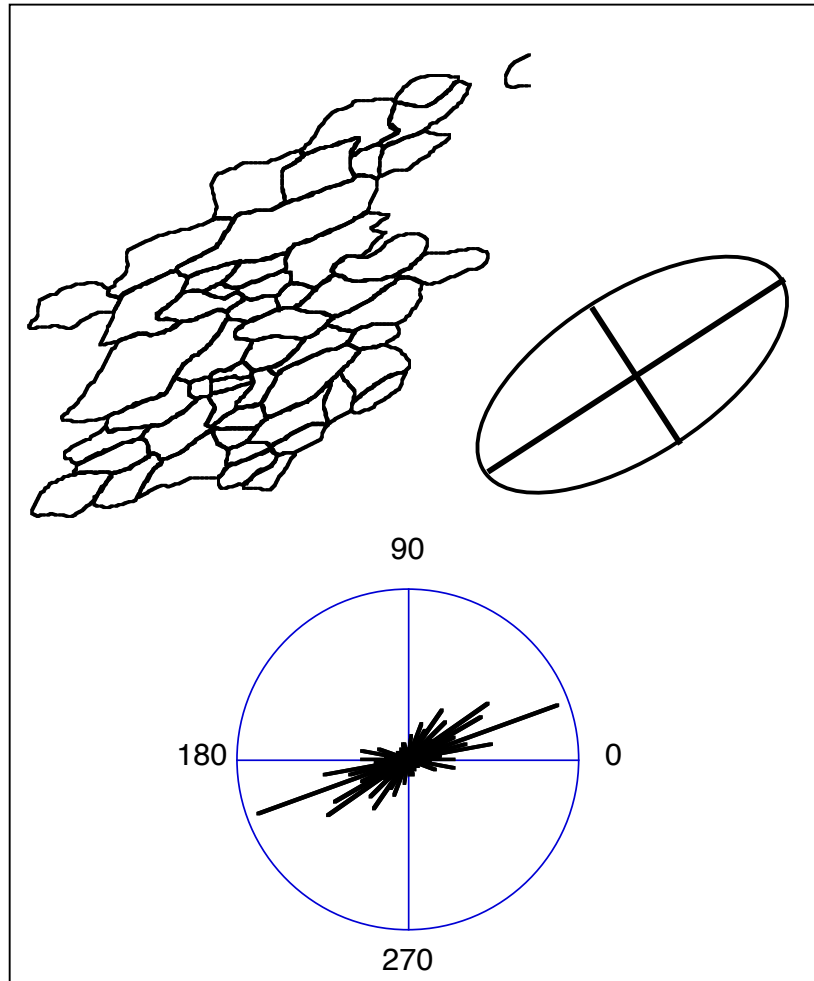
deformation mechanism map for calcite $d = 100 \mu\text{m}$

und wie erkennen wir,
ob es sich um...

dislocation glide
dislocation creep
diffusion creep

... handelt ?

microstructure of dislocation glide

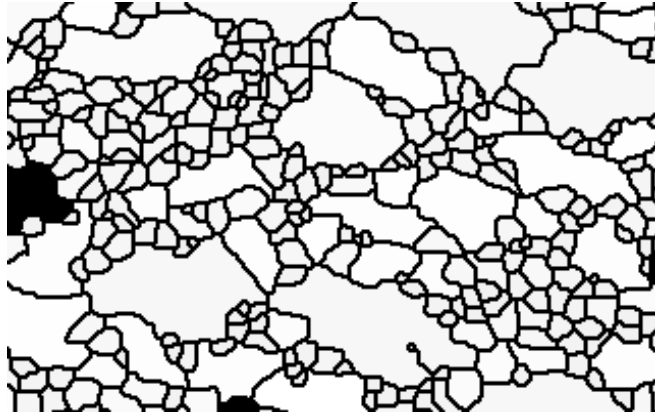


shape change = strain
strain markers,
homogeneous
deformation

typical for:

intracrystalline plasticity
dislocation glide with 5
slips systems (van Mises)

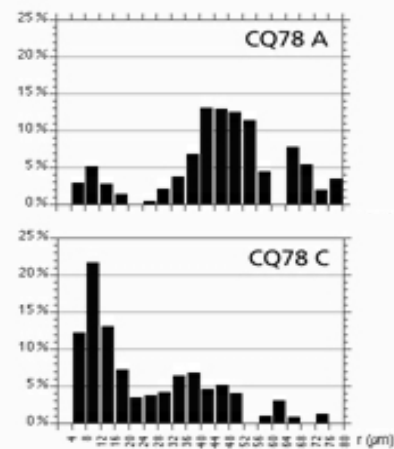
microstructure of **dislocation creep** **diffusion creep**



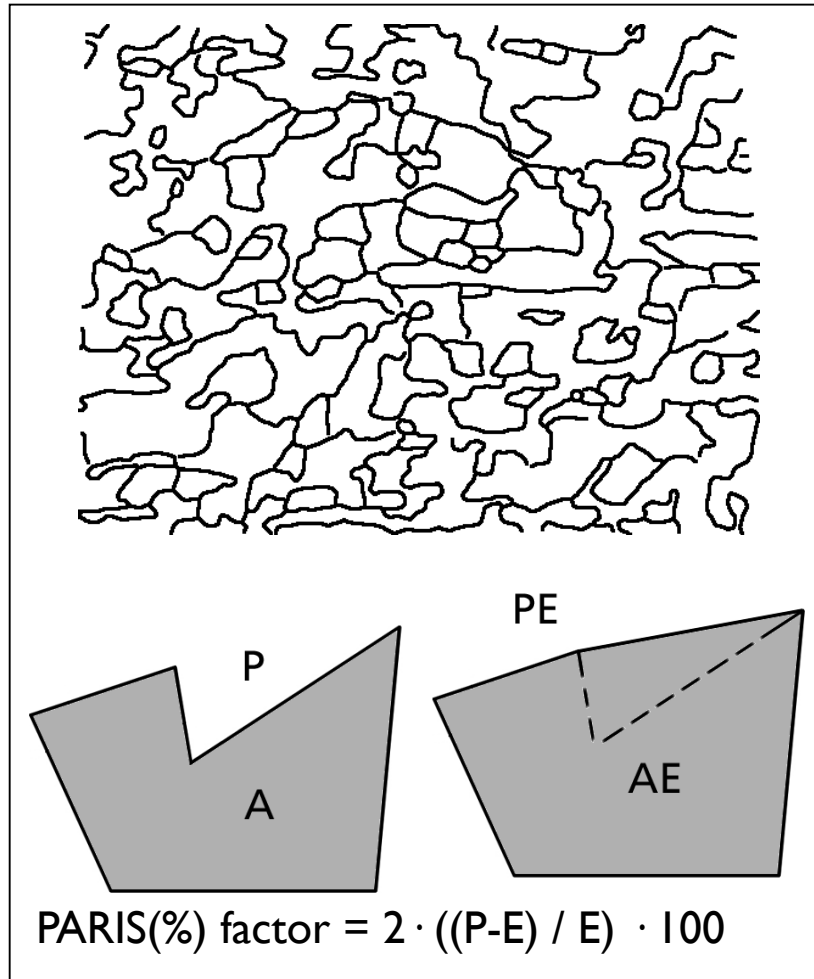
change of size
grain size distribution

typical for:

recrystallization, grain
growth, nucleation



microstructure of dislocation creep diffusion creep

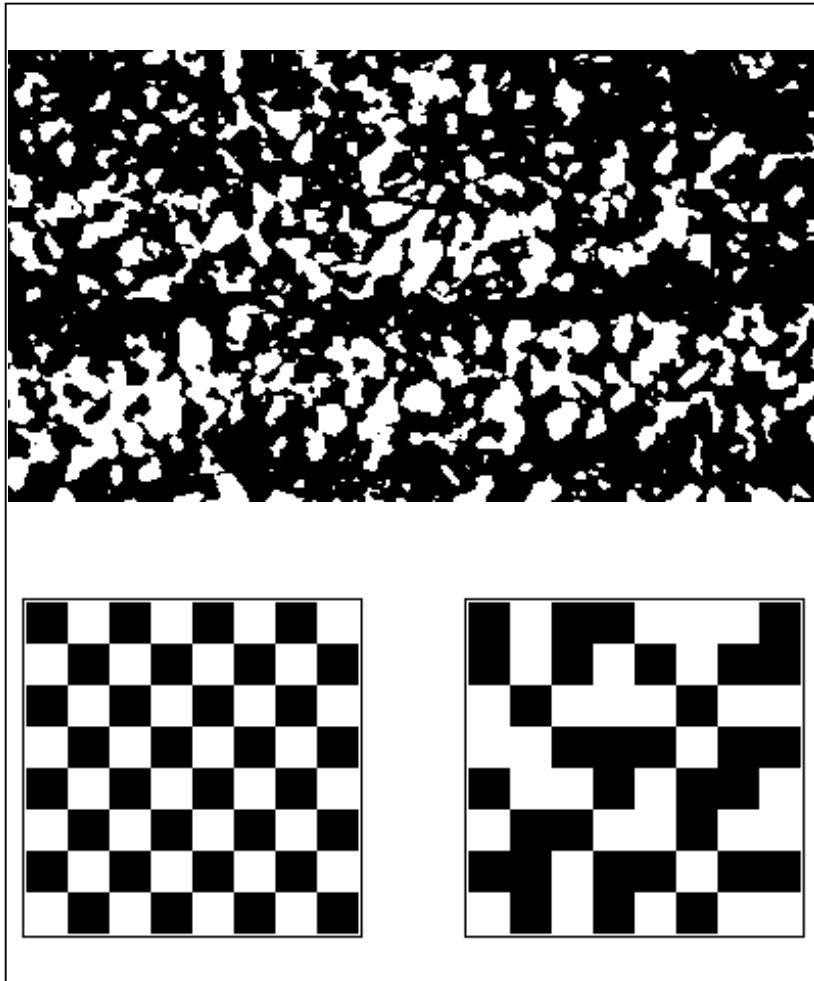


shape change \neq strain
lobate boundaries
convex-concave
angularity, etc.

typical for:

grain boundary migration
annealing, cataclasis

microstructure of diffusion creep

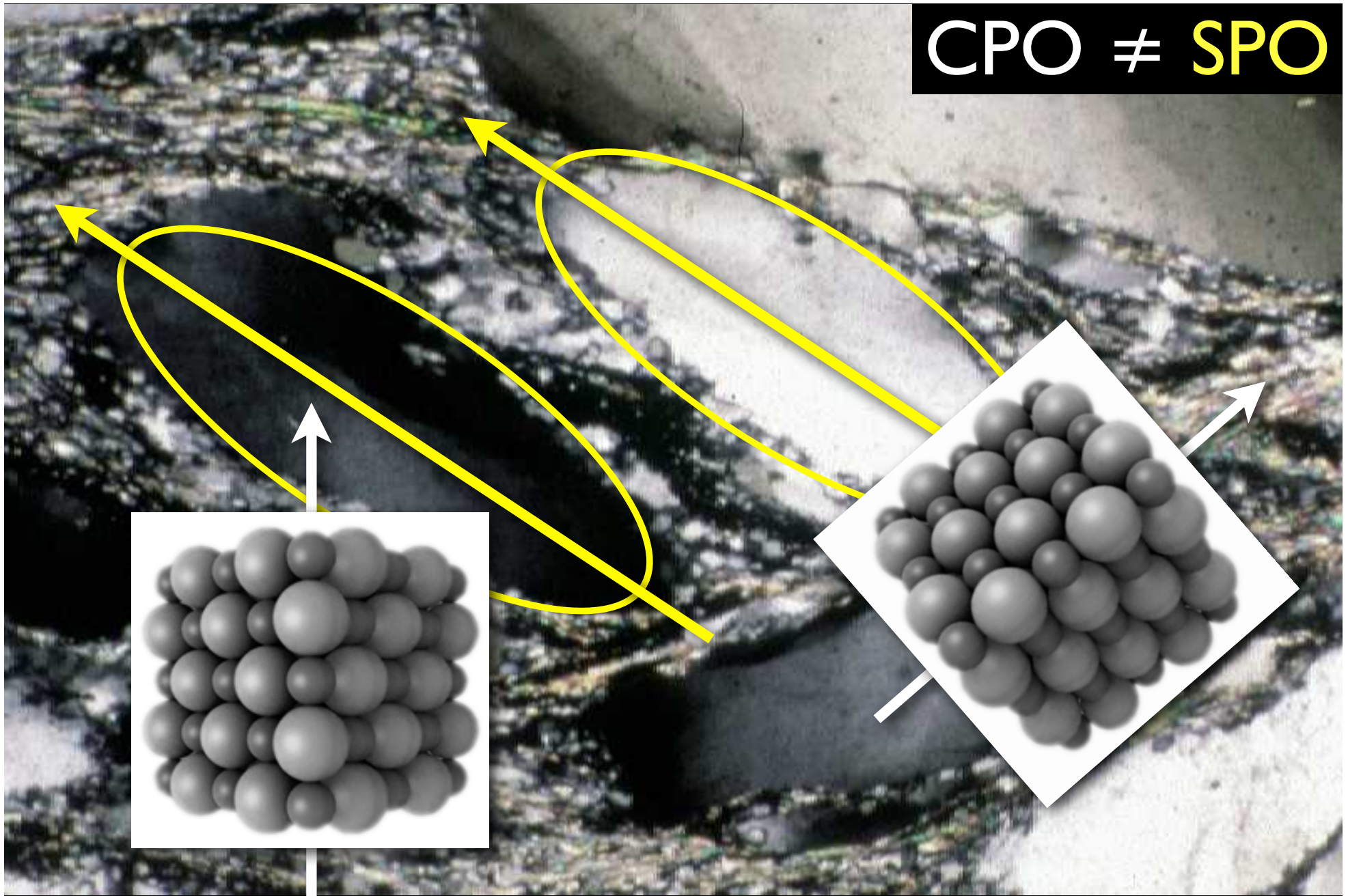


spatial distribution
grain contact frequency

typical for:

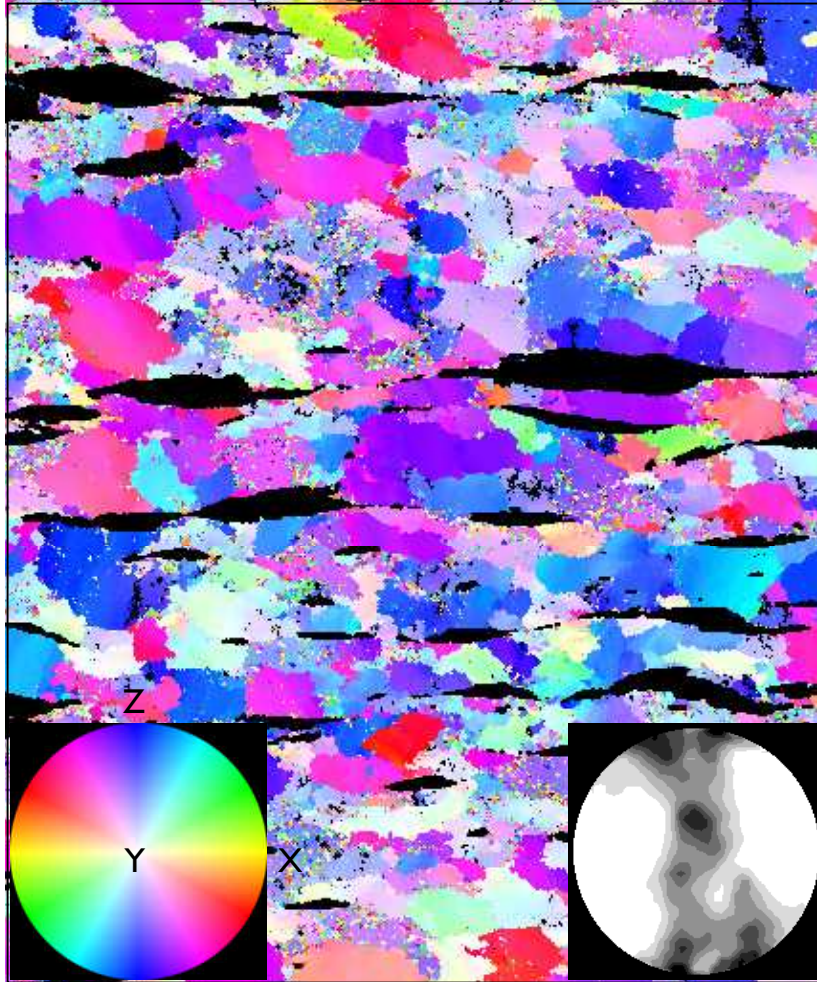
grain boundary sliding
granular flow

CPO \neq SPO



crystal orientation \neq shape orientation

CPO of dislocation creep



map of crystallographic
orientation
(orientation image)
c-axis pole figure

typical for:

dynamic recrystallization,
shear sense

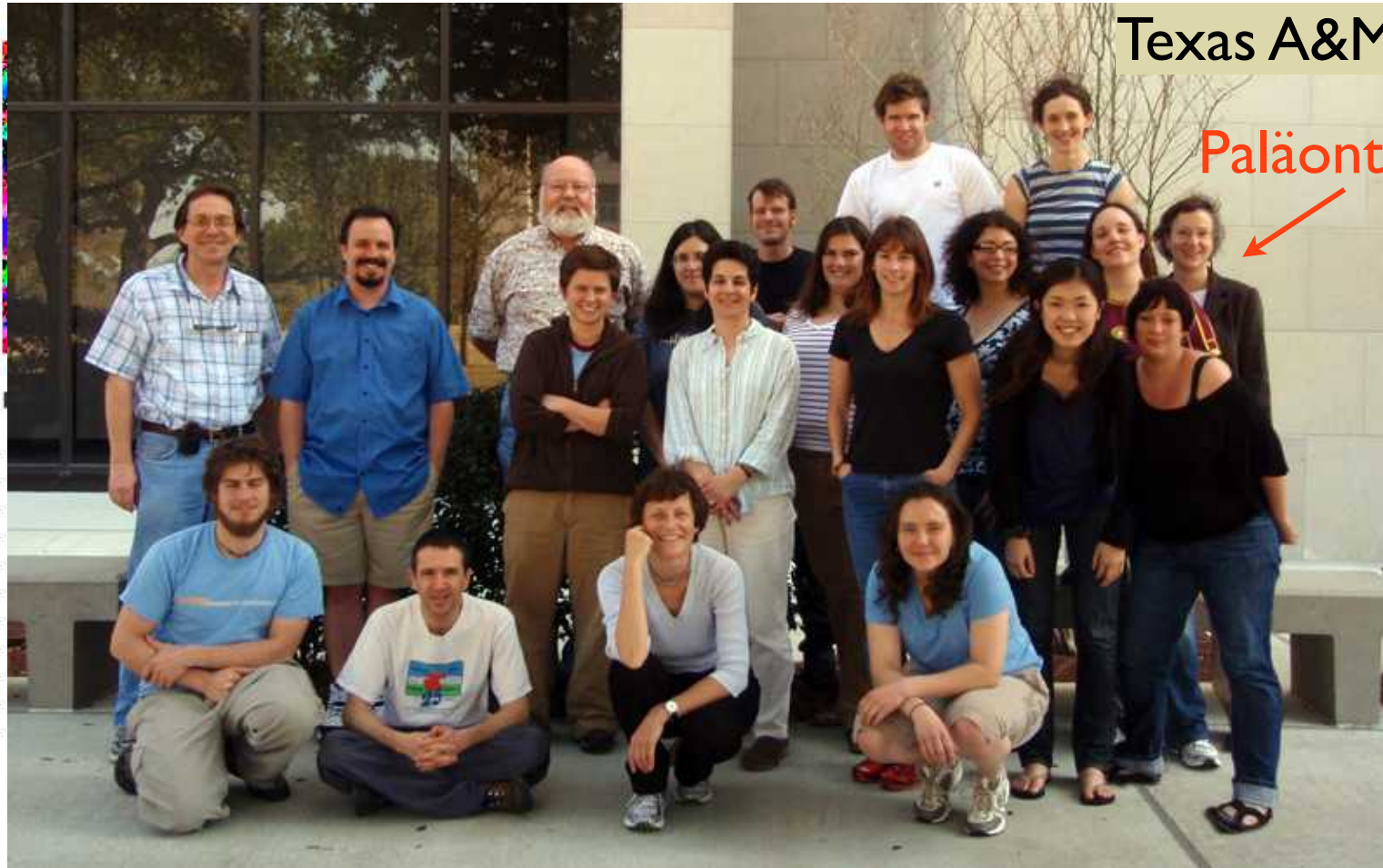
workshops

- >>> BASEL UNIVERSITY SCIENCE FACULTY (PHIL II)
- >>> BASEL UNIVERSITY LIBRARY
- >>> BASEL UNIVERSITY Vorlesungsverzeichnis
- >>> BASEL UNIVERSITY PERSSEARCH



ROCK DEFORMATION HOME

[address](#) [lectures](#) [workshops](#) [software](#) [SXM macros](#) [manuals](#) [projects](#) [publications](#)



Texas A&M

Paläontologin !!

ROCK DEFORMATION HOME

[address](#) [rock deformation lab](#) [pictures](#) [miscellaneous](#)

-impressum-

- >>> TROMSØ HARD ROCK HOMEPAGE
- >>> TROMSØ GEOLOGY HOMEPAGE
- >>> TROMSØ UNIVERSITY HOMEPAGE

>>> TROMSØ WEBGEOLOGY

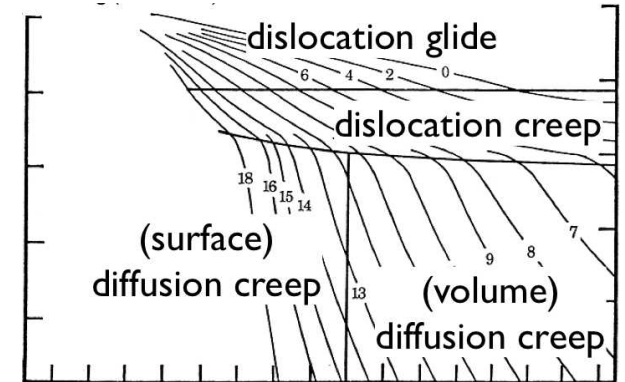


3 Dissertationen

1. dislocation glide

Anja Thust : Untersuchungen
an Quarz Einkristallen

Festigkeit der verschiedenen Gleitsysteme



2. dislocation creep - diffusion creep

Rüdiger Kilian : Deformationsmechanismen
von Quarz in polyphasen Gesteinen

3. diffusion creep

Matěj Peč : Rheologisches Verhalten von
fault rocks in Natur und Experiment.

Diss zu dislocation glide

intracrystalline deformation

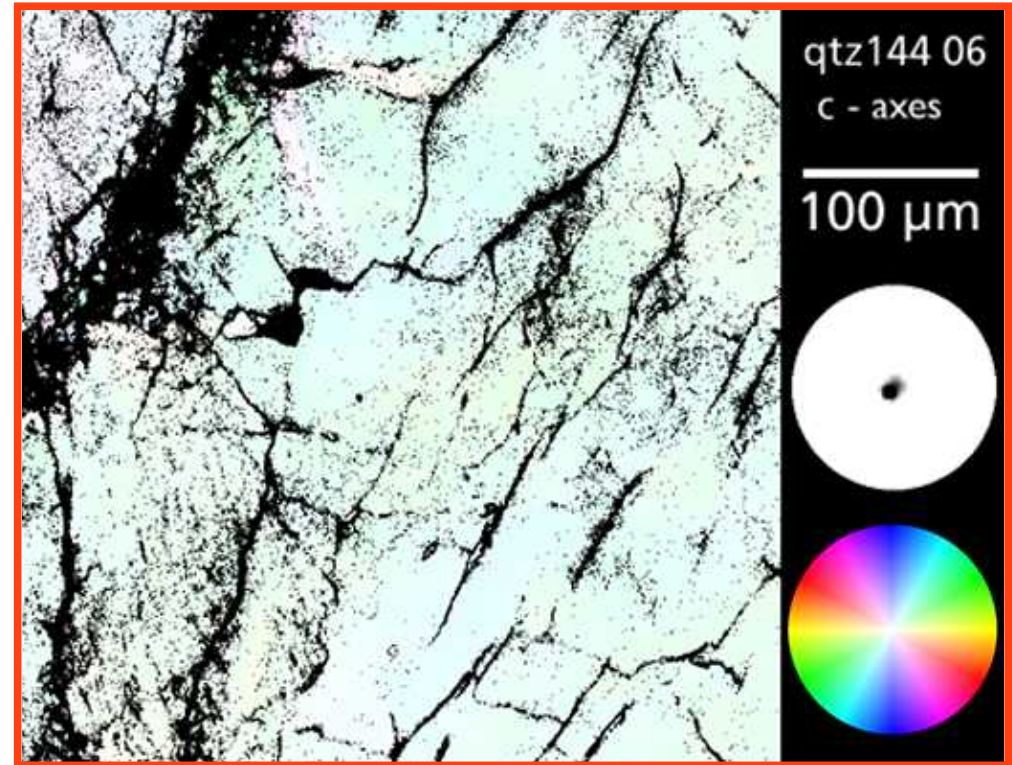
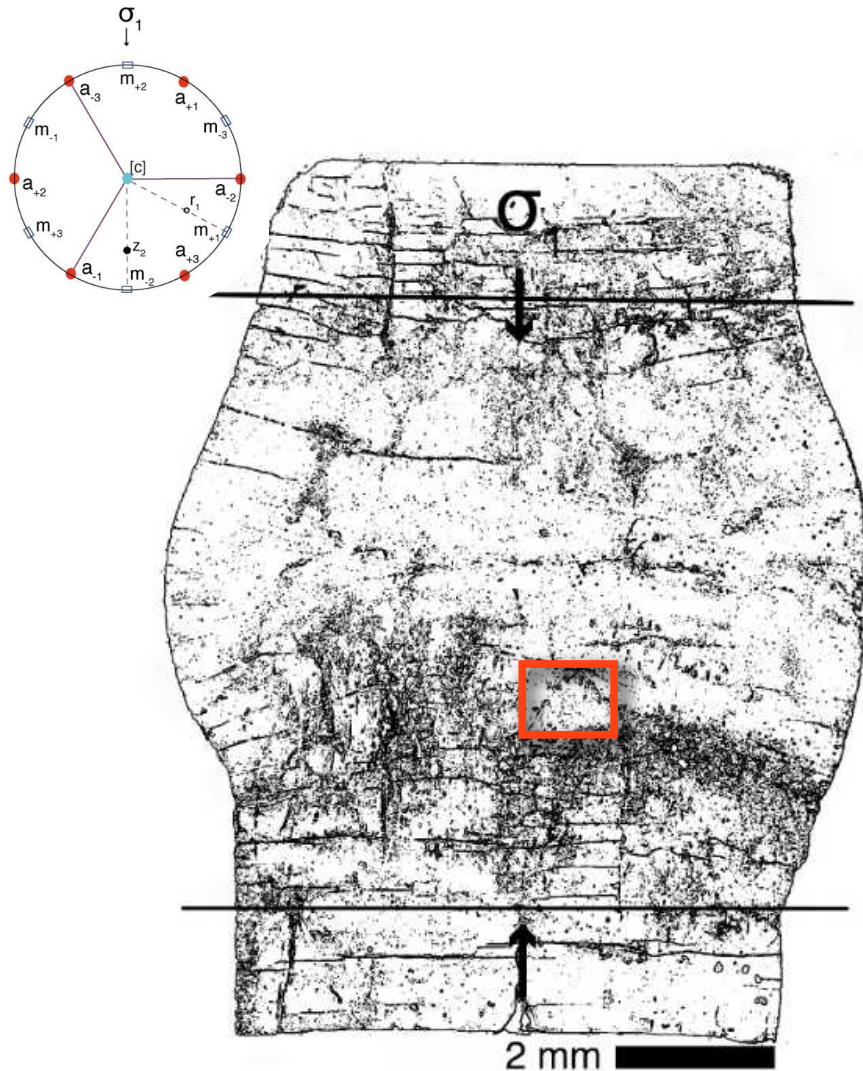
im Kristallgitter (innerhalb eines Mineralkornes) ist dislocation glide "unsichtbar"

das Gitter ist selbstidentisch
nur die äussere Form ändert

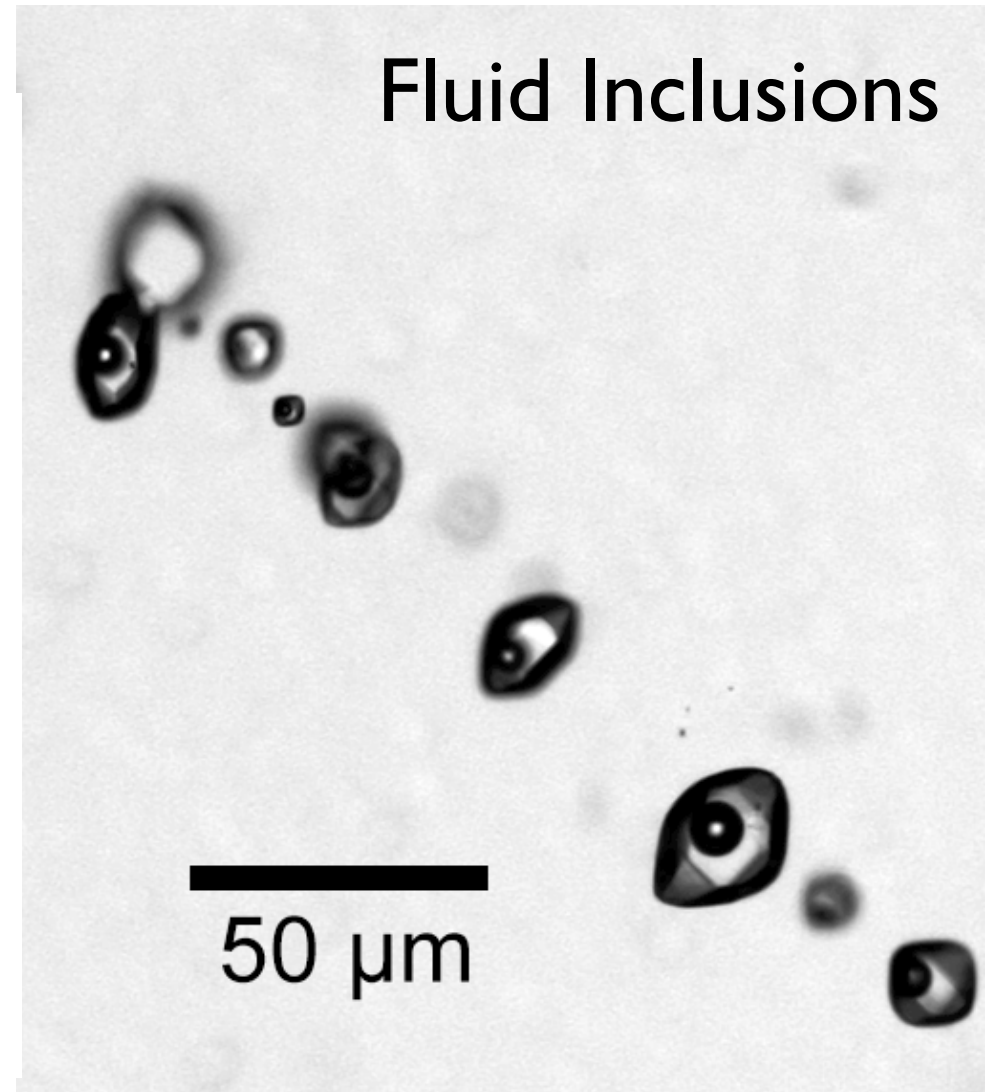
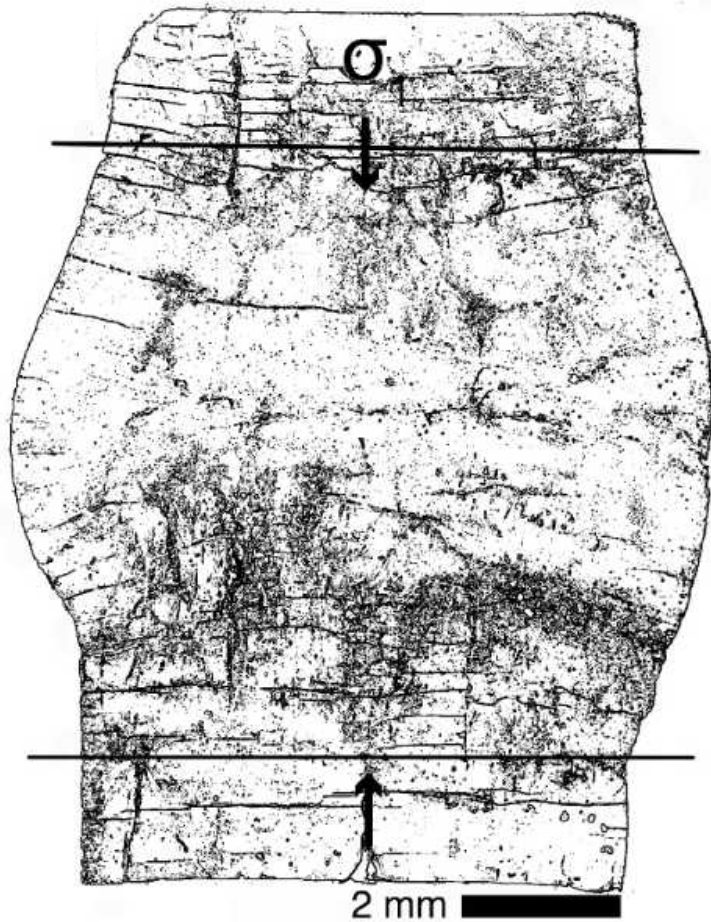
2 mm

2 mm

Anja Thust : Quarz - Einkristalle



Anja Thust : Quarz - Einkristalle



Diss zu dislocation creep

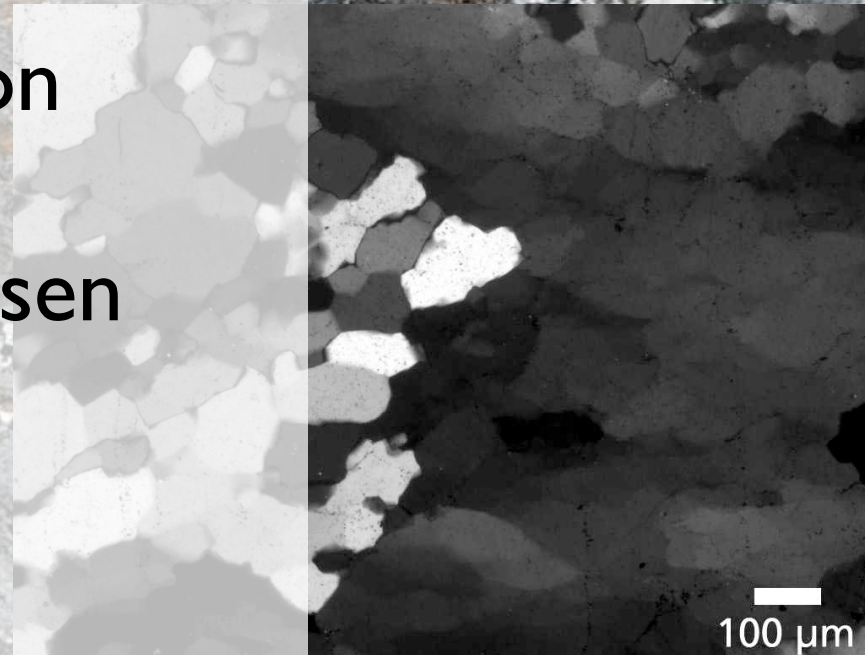
intra- intercrystalline deformation

dynamische Rekristallisation

Übergang von 2 rheol. Phasen
(Quarz & Matrix)

zu homogenem chem.

Phasengemisch = 1 rheol. Phase

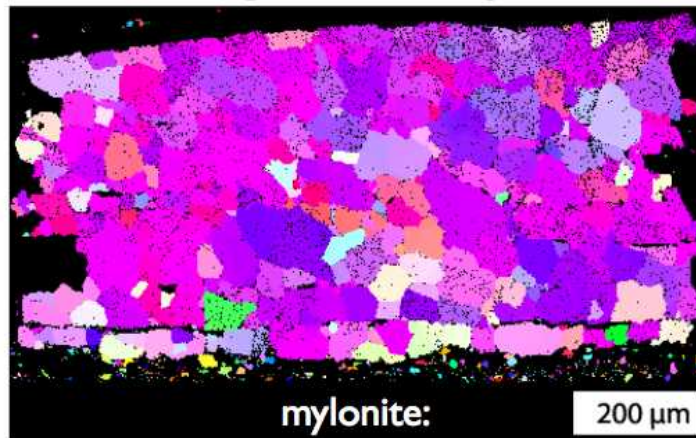


4 mm

Rüdiger Kilian : Quarz in Polyphasen

dislocation
creep

Quartz grain shape:

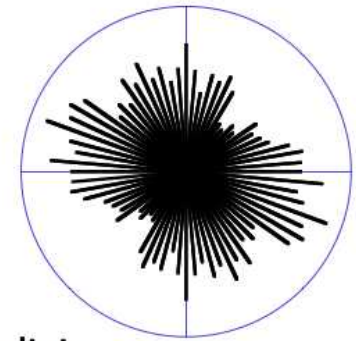
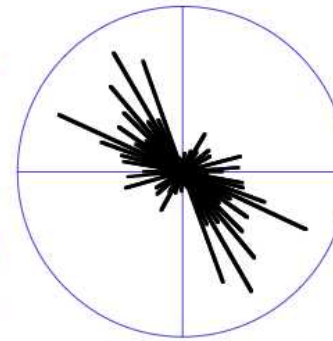


diffusion
creep

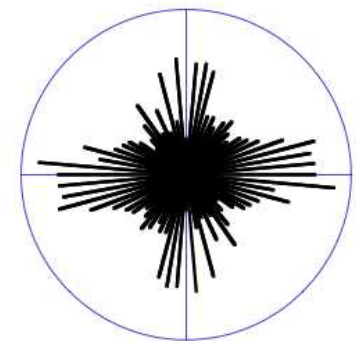
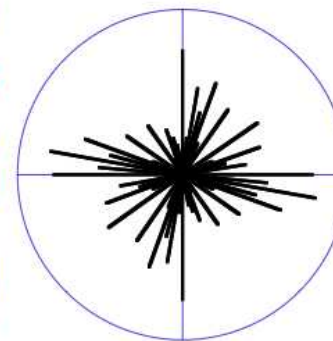


paror
(particle orientation)

surfor
(surface orientation)



monoclinic

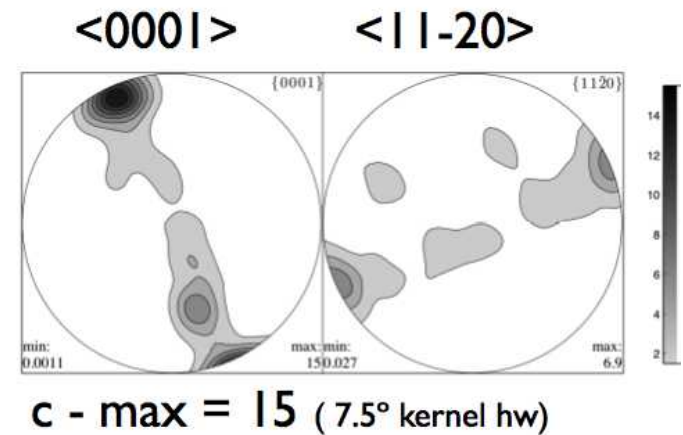
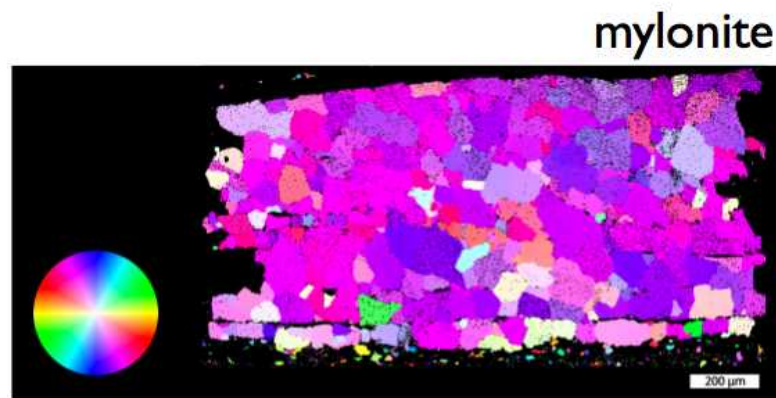


~ orthorhombic

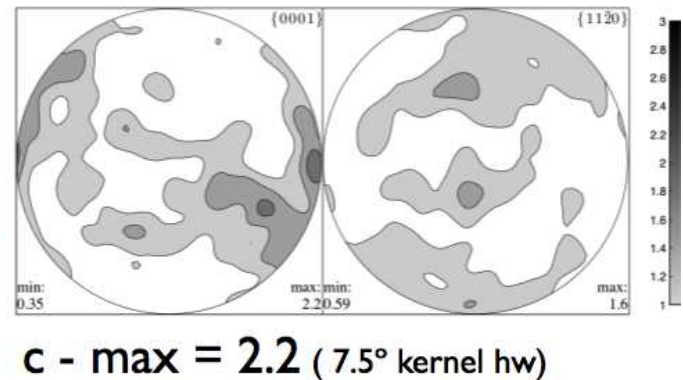
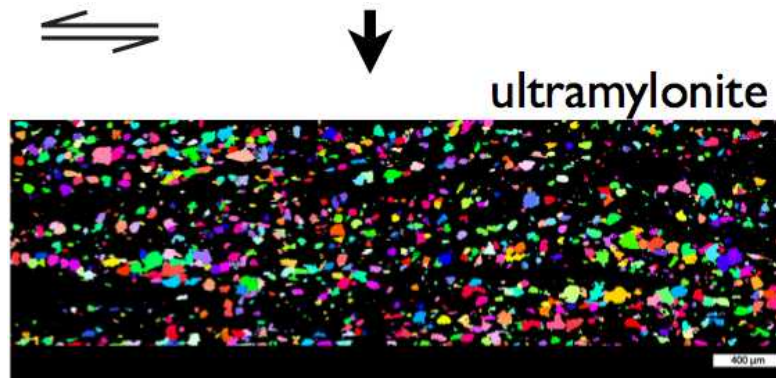
Rüdiger Kilian : Quarz in Polyphasen

Quartz texture:

dislocation
creep



diffusion
creep



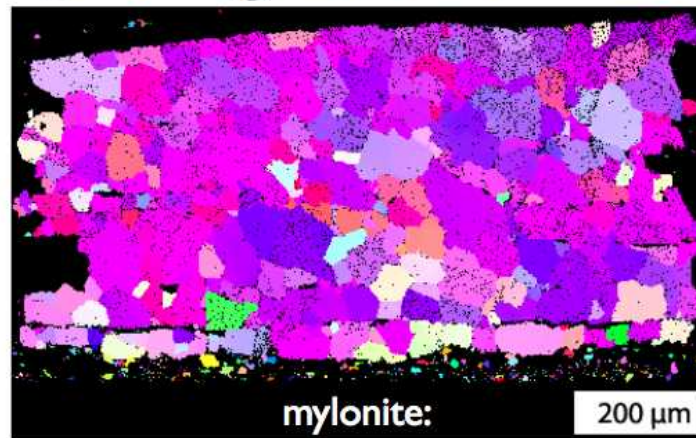
EBSD data with CIP c-axes orientation colour coding

calculated with MTEX (Hielscher & Schaeben, 2008)

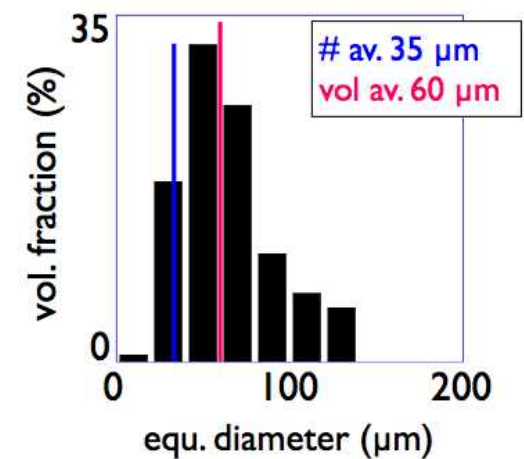
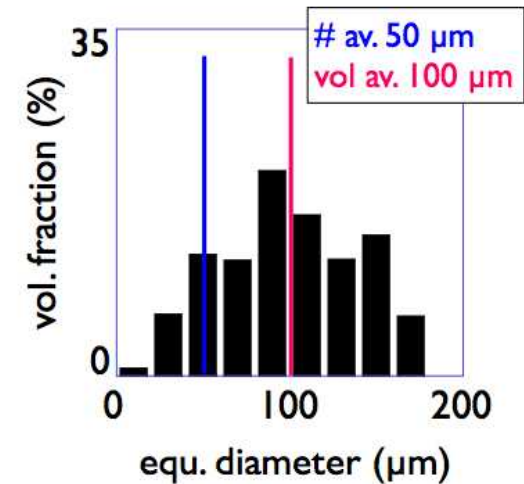
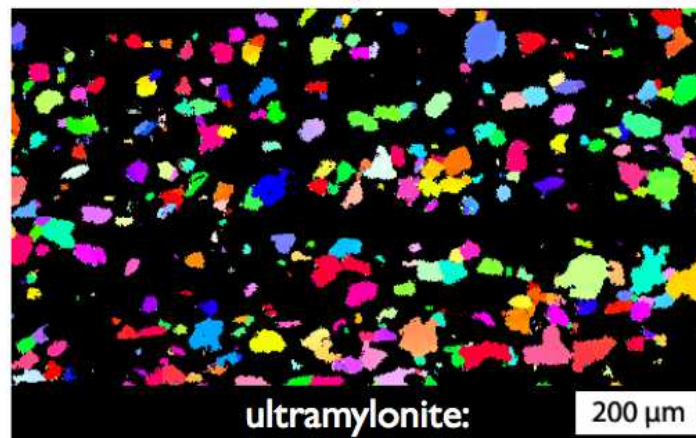
Rüdiger Kilian : Quarz in Polyphasen

dislocation
creep

Quartz grain size:

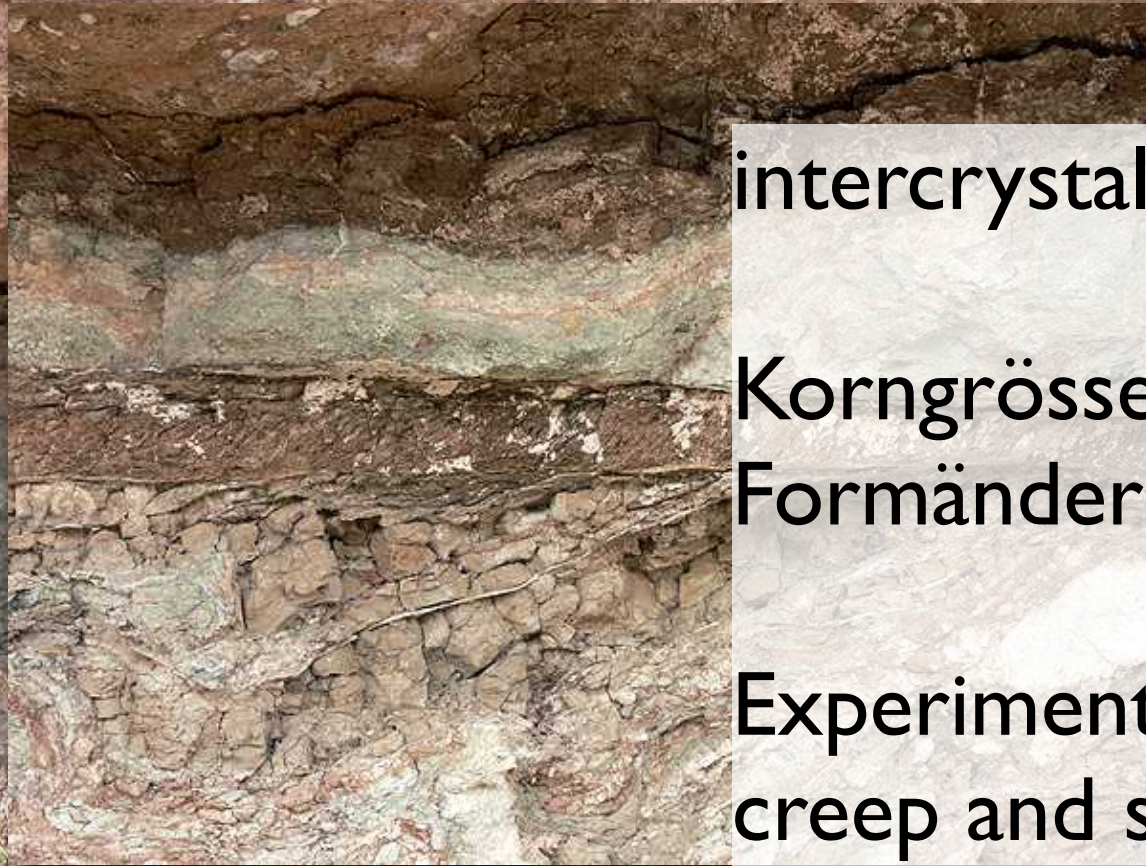


diffusion
creep



Diss zu diffusion creep

Matěj Peč : Fault Rocks



intercrystalline deformation

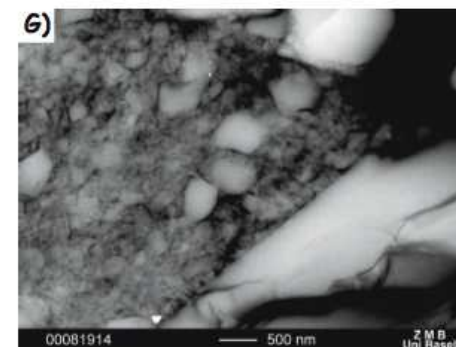
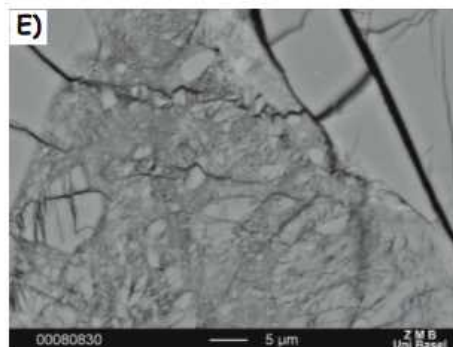
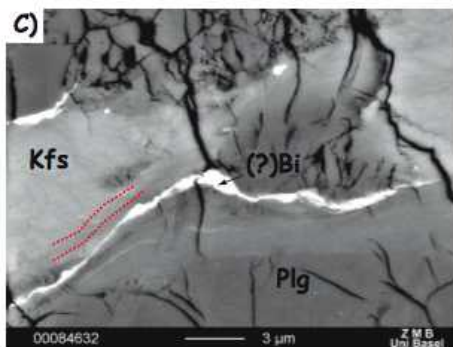
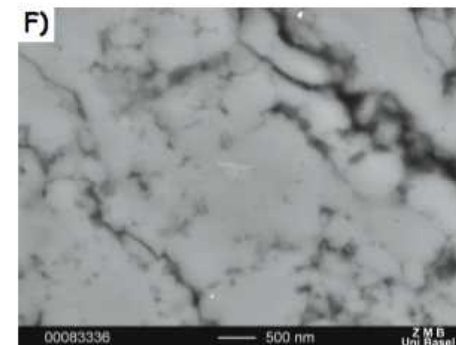
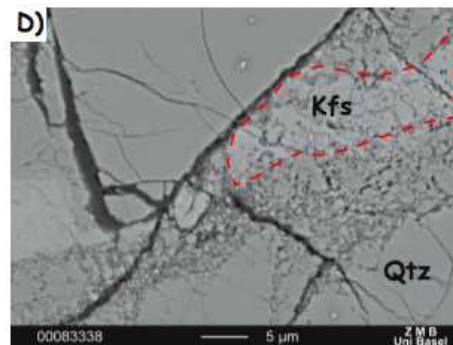
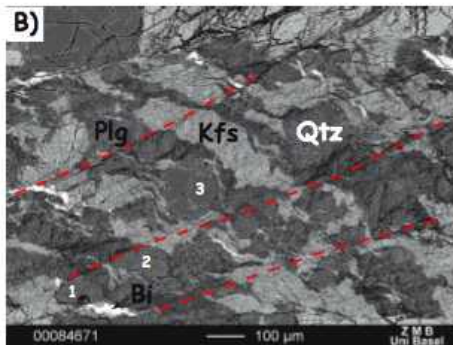
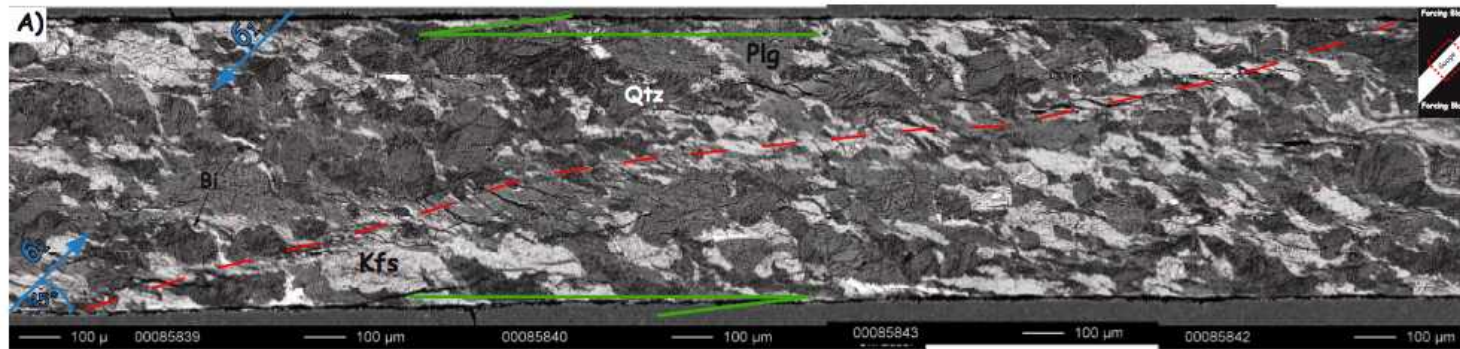
Korngrößenabhängigkeit
Formänderung

Experiment - Natur:
creep and slip and healing
strength of faults

diffusion creep

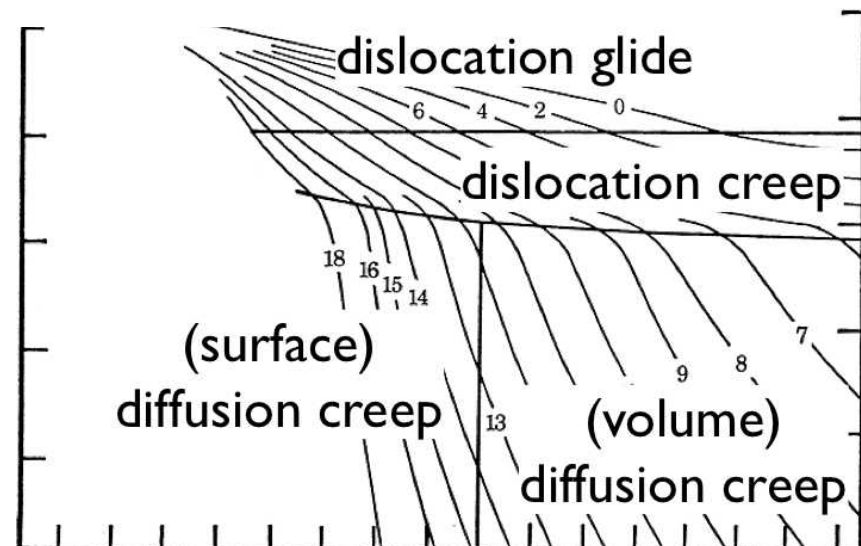


Matěj Peč : Fault Rocks



Moral und Nutzanwendung

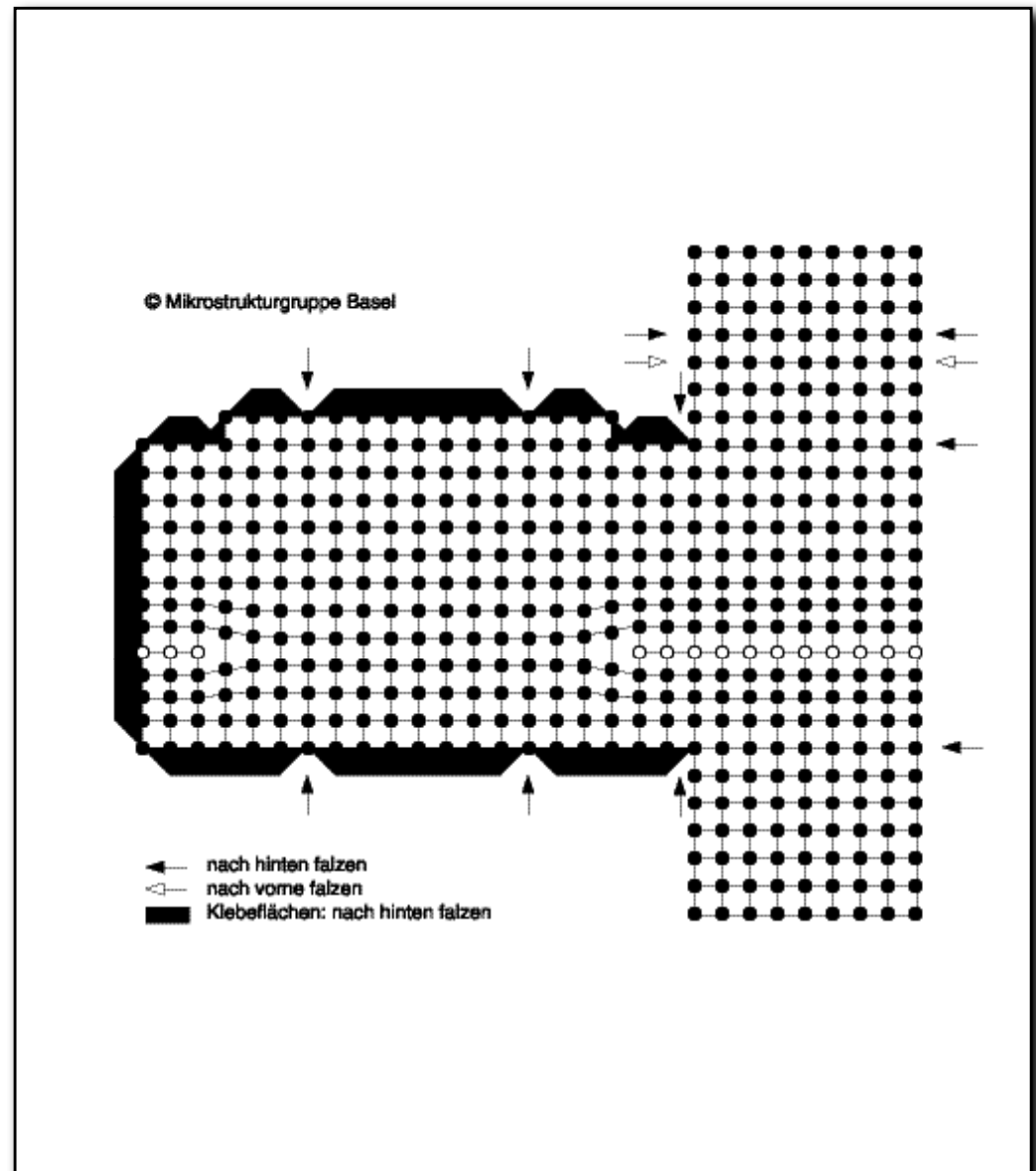
wie aus vielen schnellen
Versetzungsen eine grosse
Langsamkeit entsteht



Moral und Nutzenanwendung

- es braucht Fehlstellen
- noch besser Korngrenzen
- noch besser Zweit-/Drittphasen
- Wasser hilft auch

zum Schluss:
slow dislocations
... zum Heimnehmen



the end

